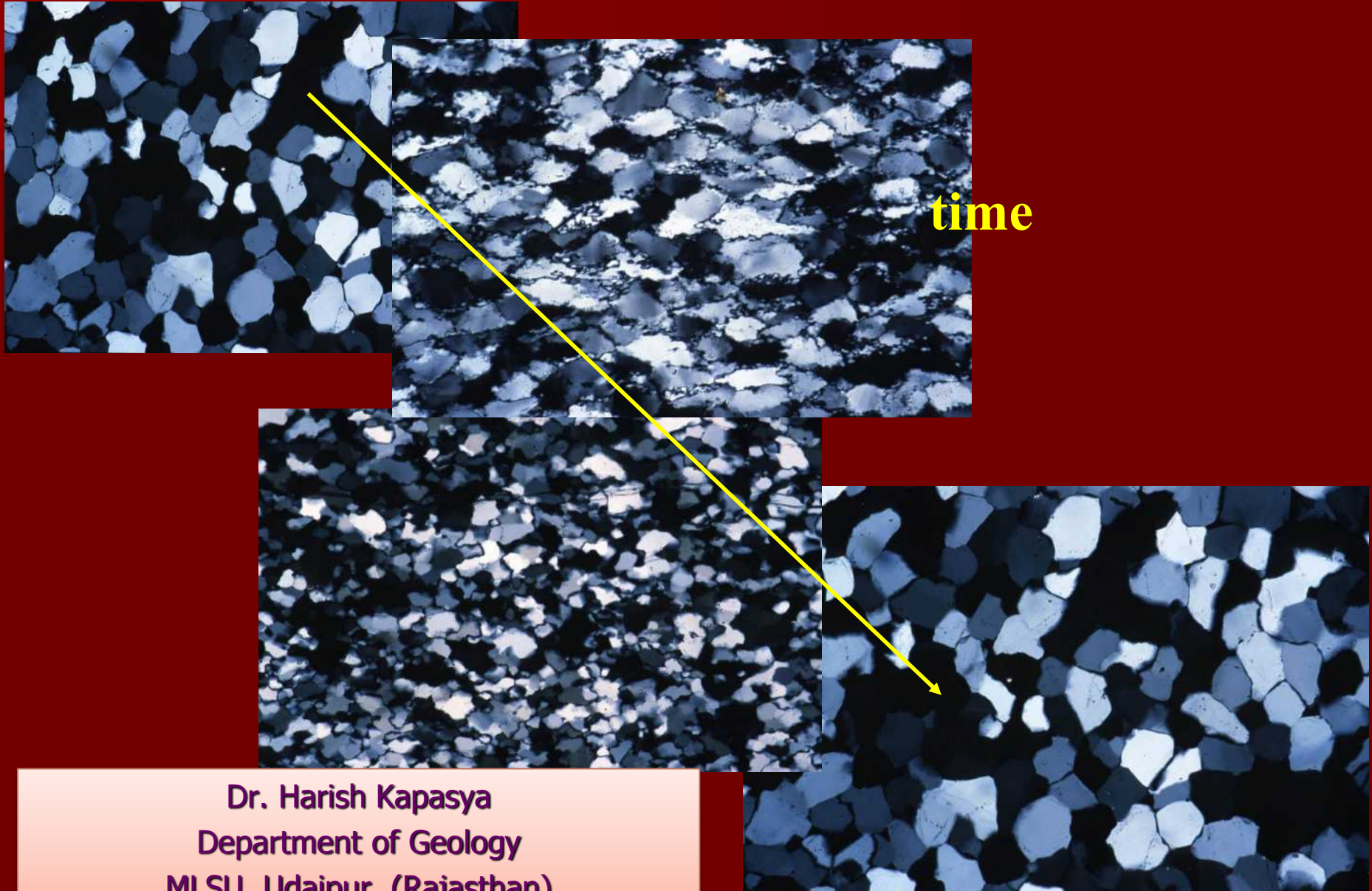
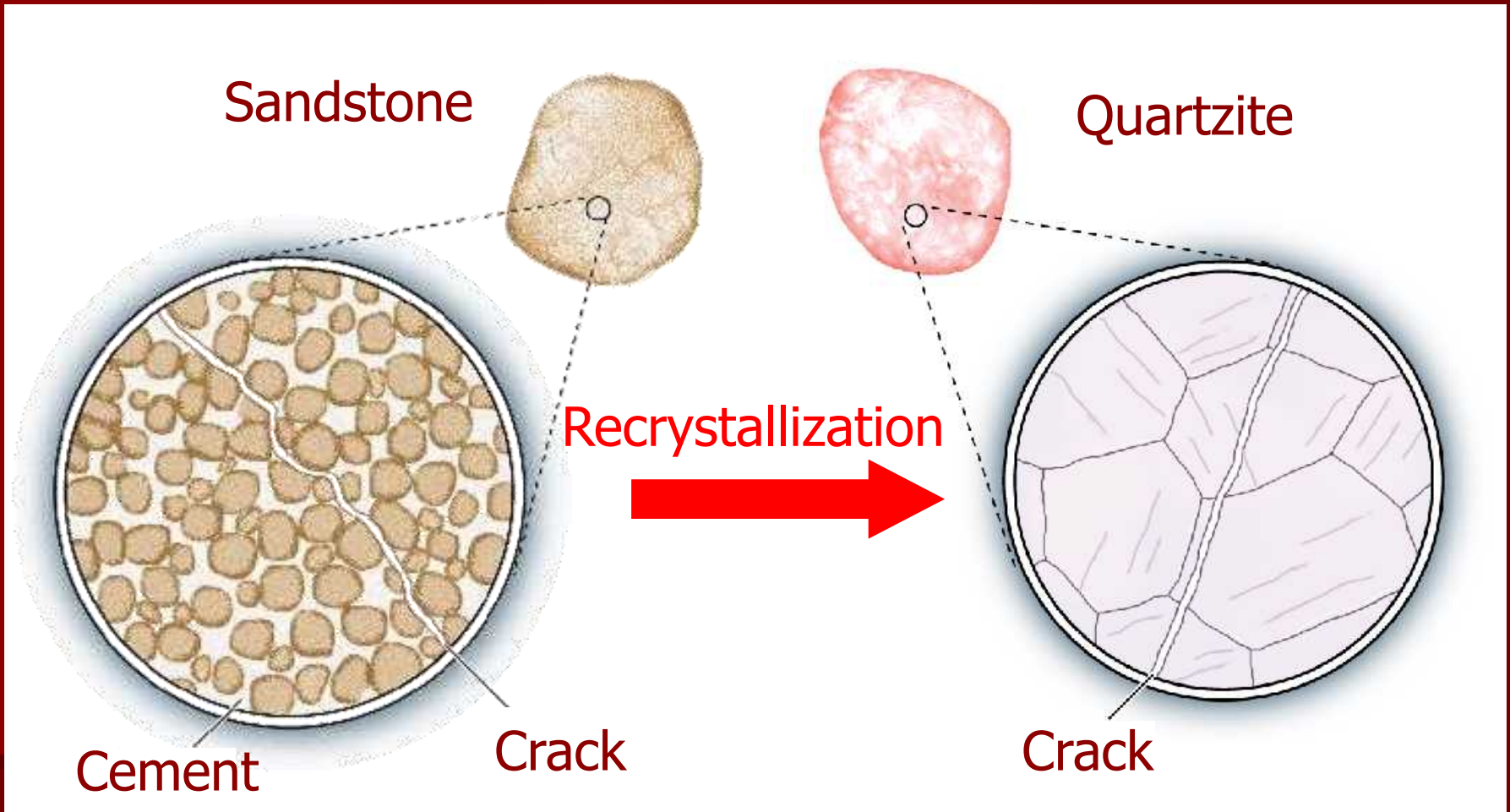


# Relationships between deformation and metamorphism

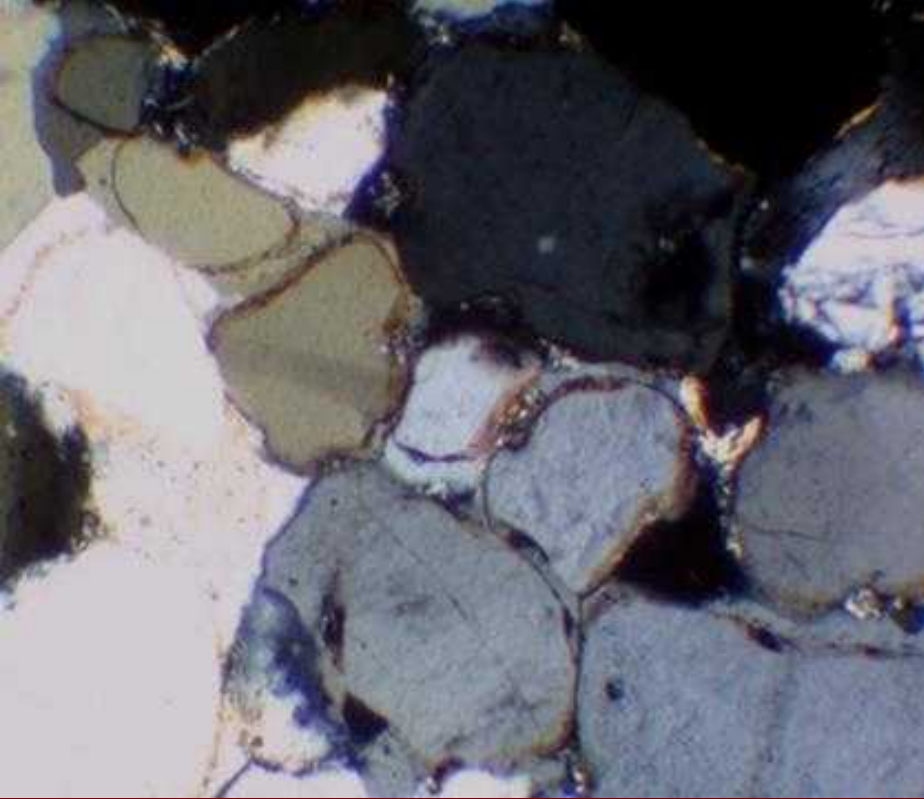


Dr. Harish Kapasya  
Department of Geology  
MLSU, Udaipur, (Rajasthan)

# Recrystallization: Minimization of Interfacial Free Energy by Minimizing Surface Area





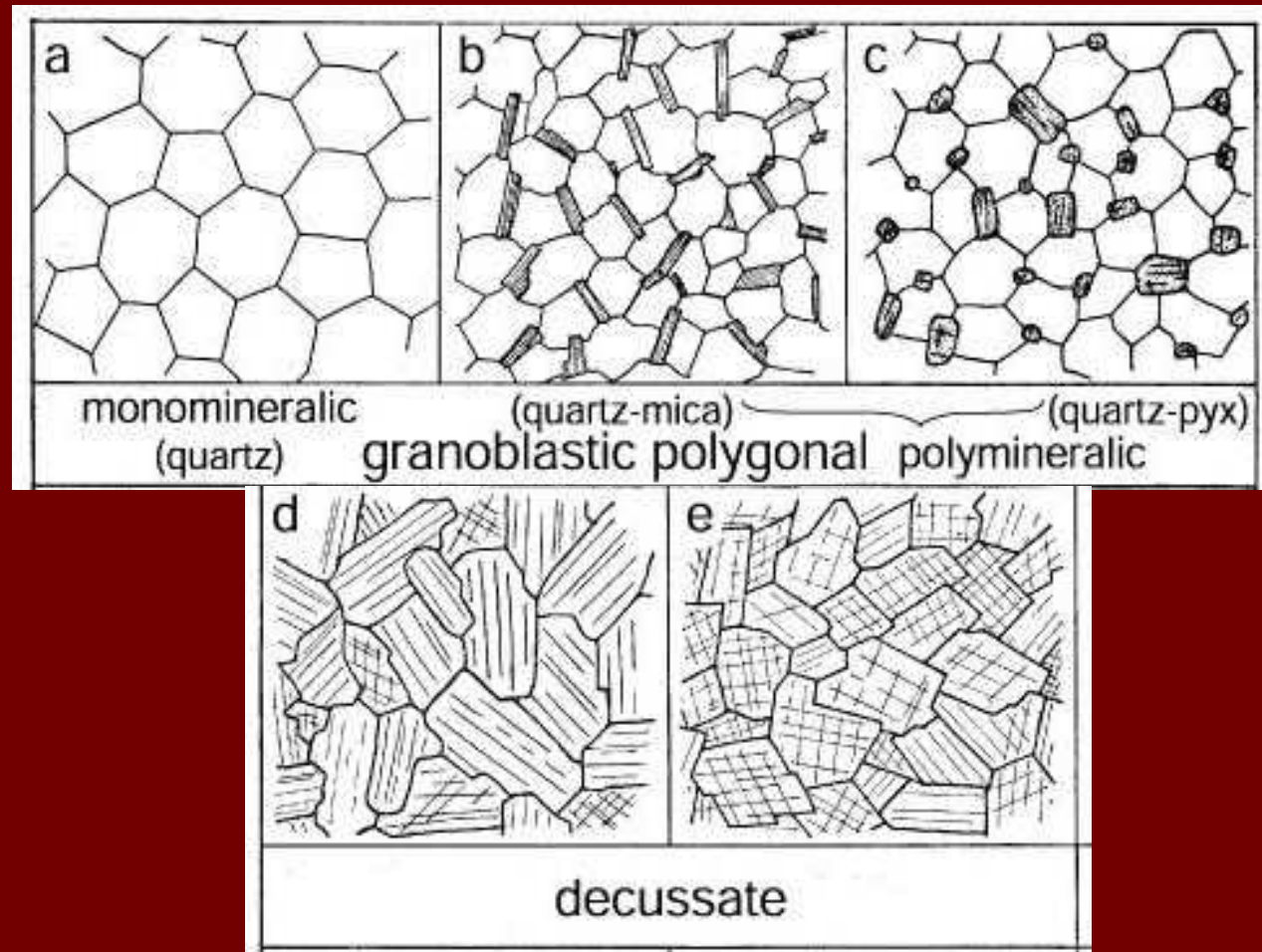


Sandstone  
texture



Quartzite  
texture

# Granoblastic/Decussate Textures (Characteristic of Contact or Thermal Metamorphism)



## **FOLD GENERATION:**

folds that formed at approximately the same time  
*...commonly find several fold generations in an area...*

each generation labeled by letter F (fold) and number;  
number reflects *relative* order of formation...

$F_1$  (first);  $F_2$  (second);  $F_3$  (third), etc.

several fold generations may fold during an orogenic phase

## ***OROGENIC PHASE:***

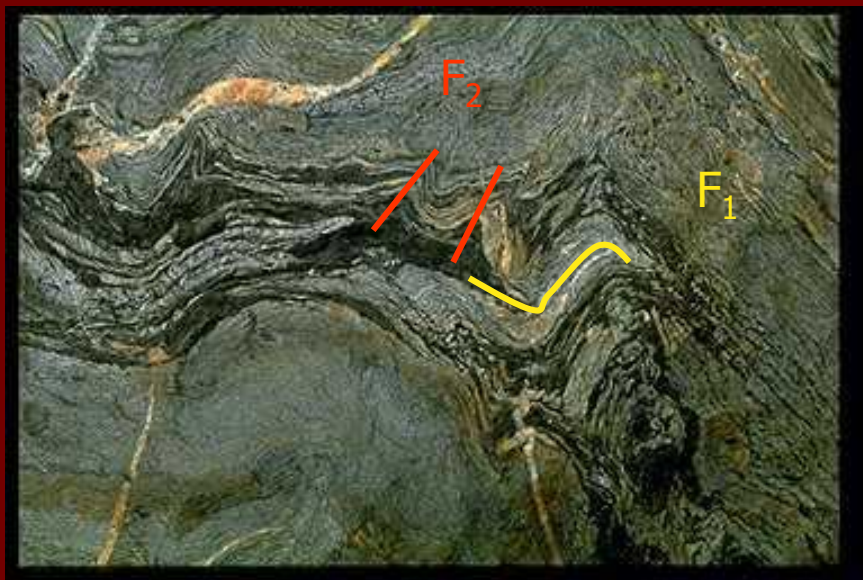
*mountain-building event;*  
given letter D (deformation) ( $D_1$ ,  $D_2$ ,  $D_3$ , etc.)



to distinguish folds of different generations  
use **principle of superposed folding**...

...later generations are superposed on, or overprint, earlier ones  
*fold superposition creates specific patterns --requires careful  
geometric analysis*

**simple rule: a superposed fold is younger than structure it folds**



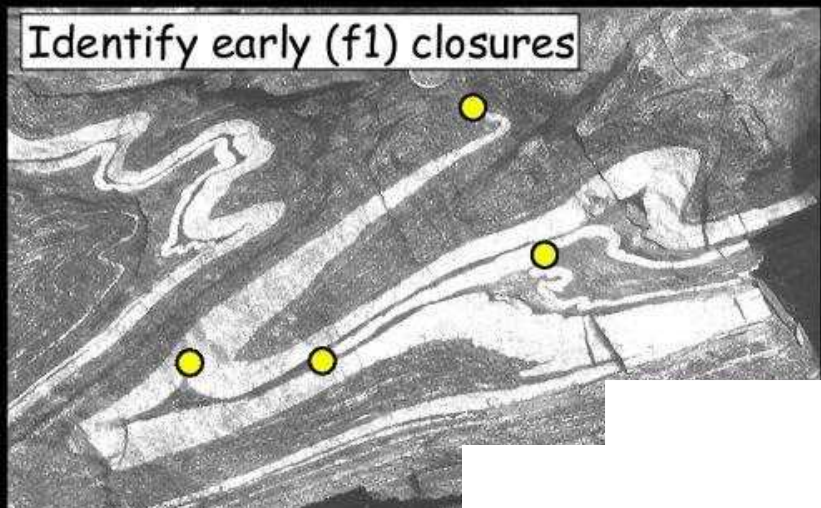
**examples of superposed folding**

*note axial surface of early fold is folded, later axial surface is not*

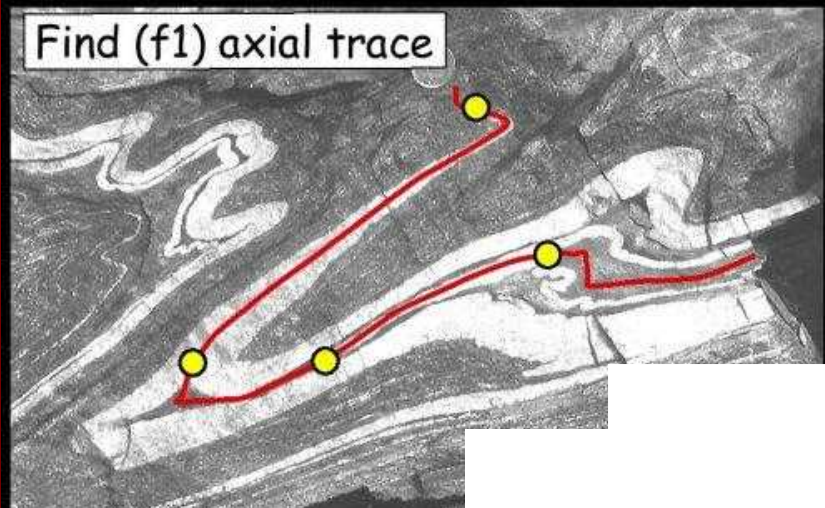
*we will examine right image more closely...*

# PROCEDURE FOR EXAMINING SUPERPOSED FOLDS

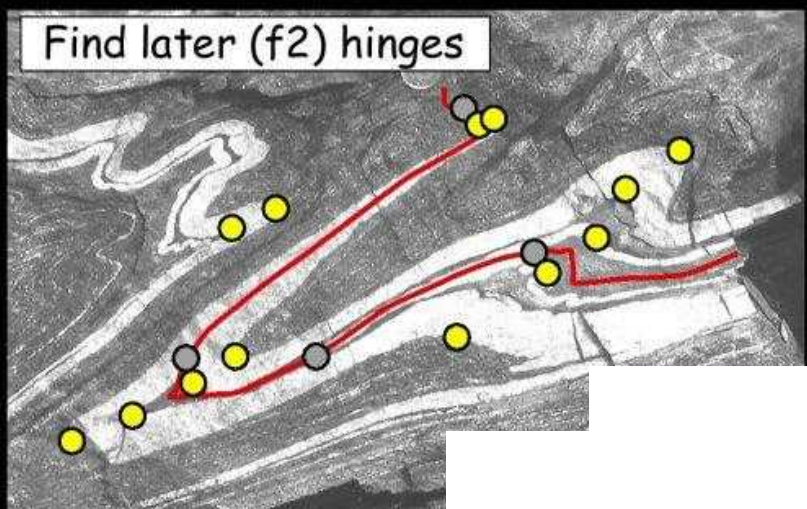
Identify early (f1) closures



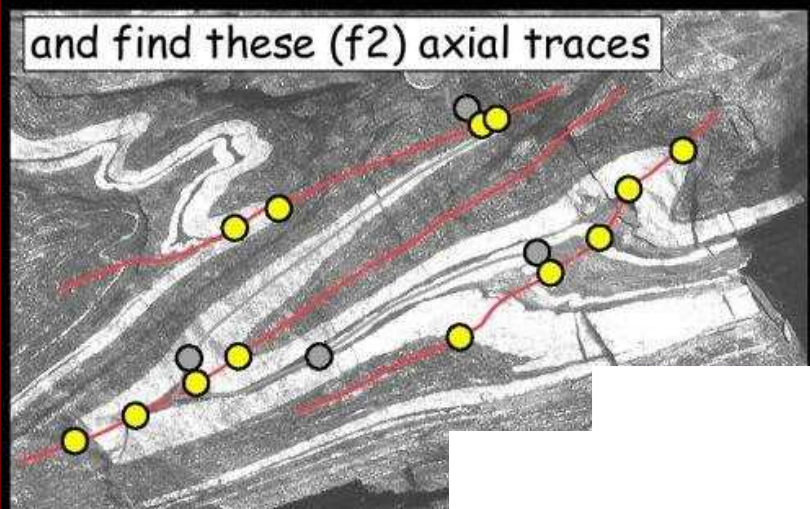
Find (f1) axial trace



Find later (f2) hinges

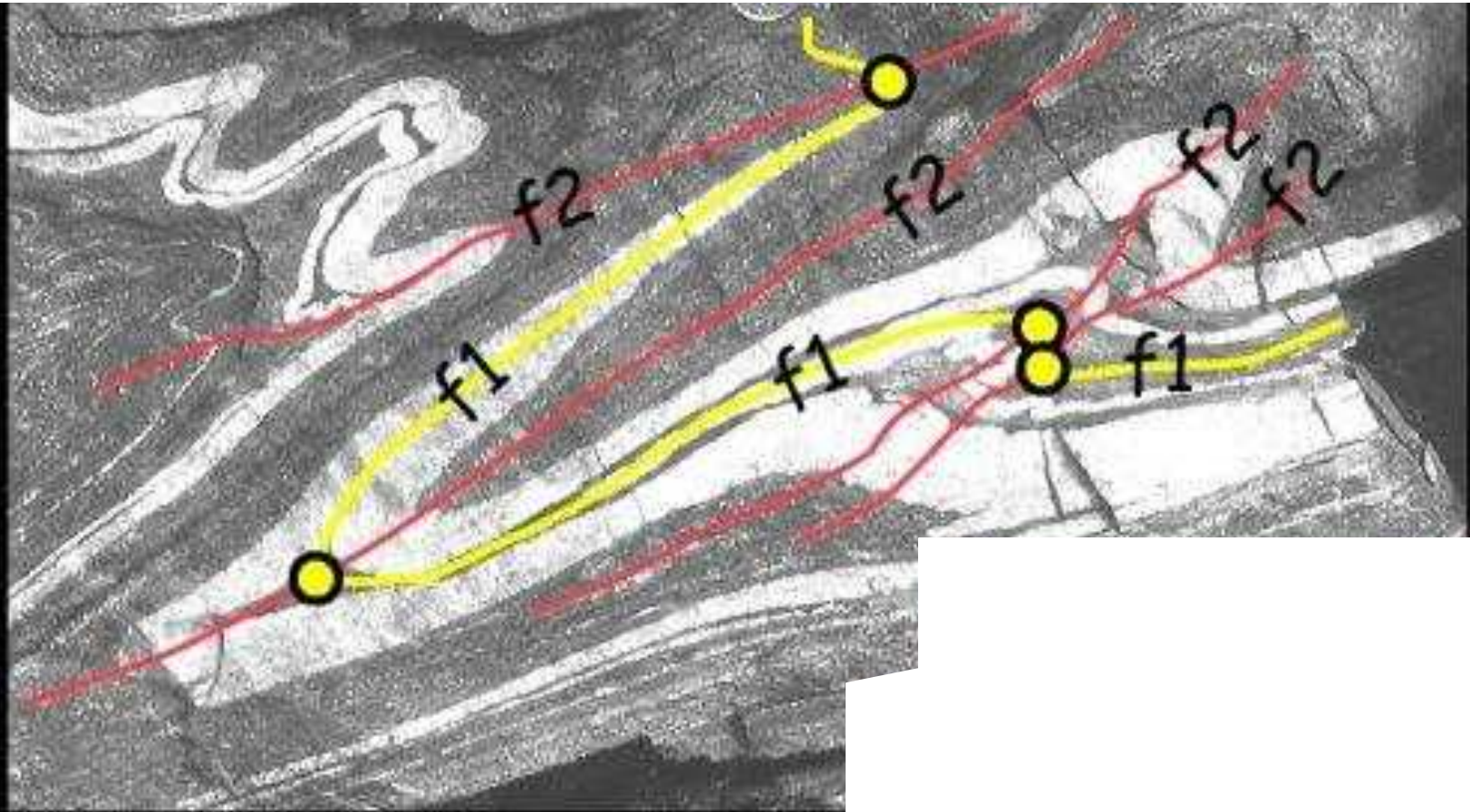


and find these (f2) axial traces





both folds identified...



**$F_1$  axial surface is folded;  $F_2$  is not. rule of superposition shows that  $F_2$  is younger than  $F_1$**



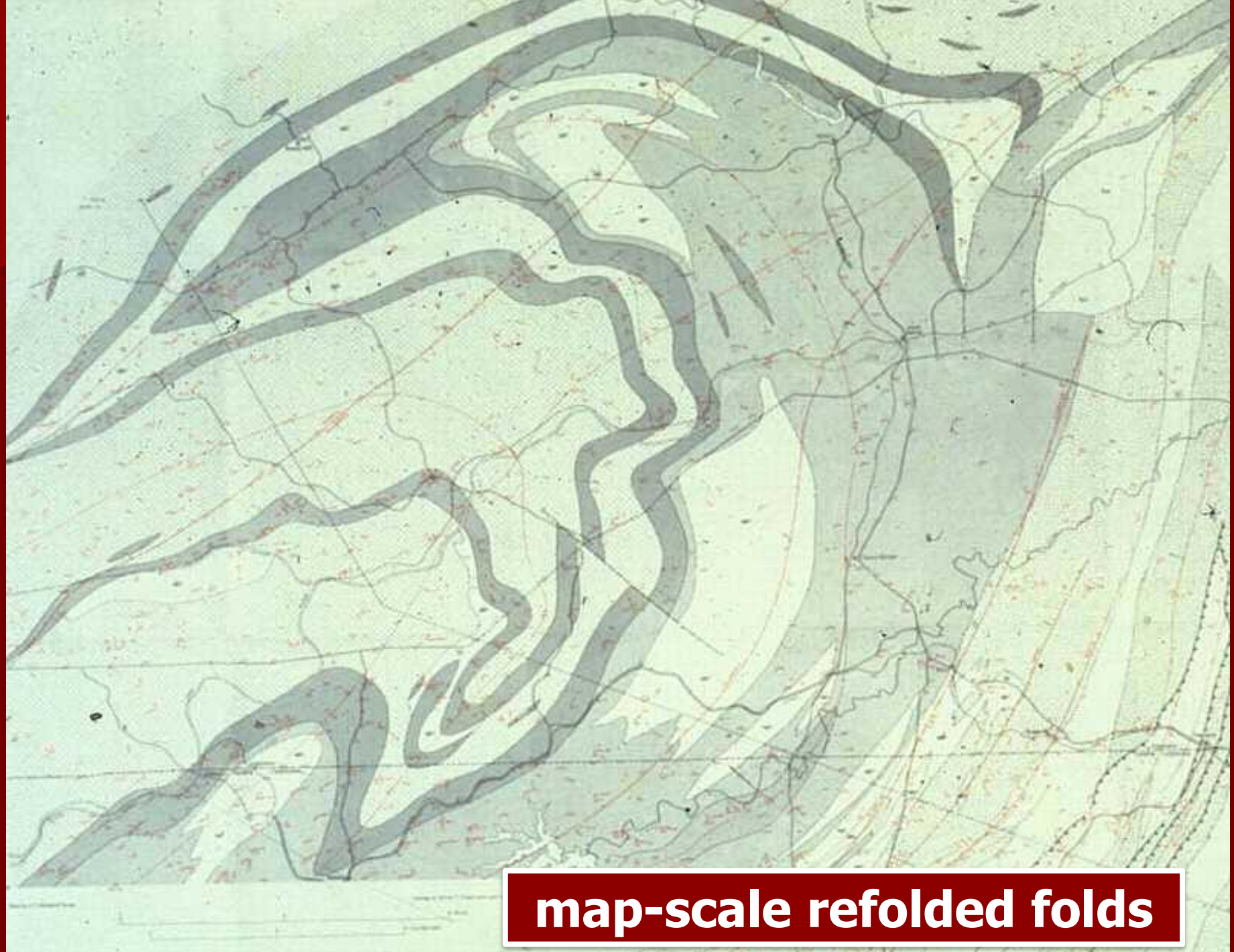




more folded folds...



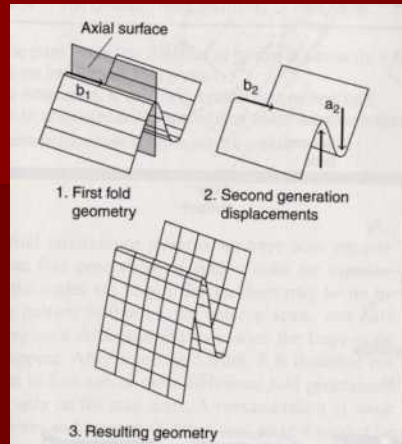




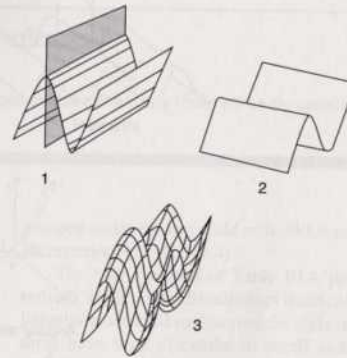
**map-scale refolded folds**

## 4 basic endmember interference patterns are seen from superposition of upright $F_2$ folds on $F_1$ folds of variable orientation

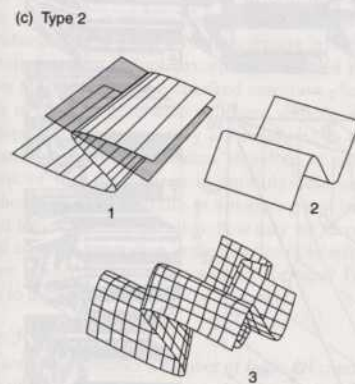
type 0



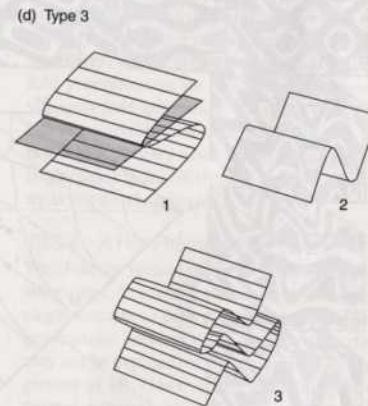
type 1



type 2



type 3



- type 0: cannot see as interference type; axial surfaces parallel
- type 1: "dome-and-basin" structure; egg-carton; axial surfaces normal
- type 2: most difficult to visualize; "mushroom"
- type 3: "refolded folds" (all types are refolded)





type 1 interference: egg-carton;  
dome-and-basin



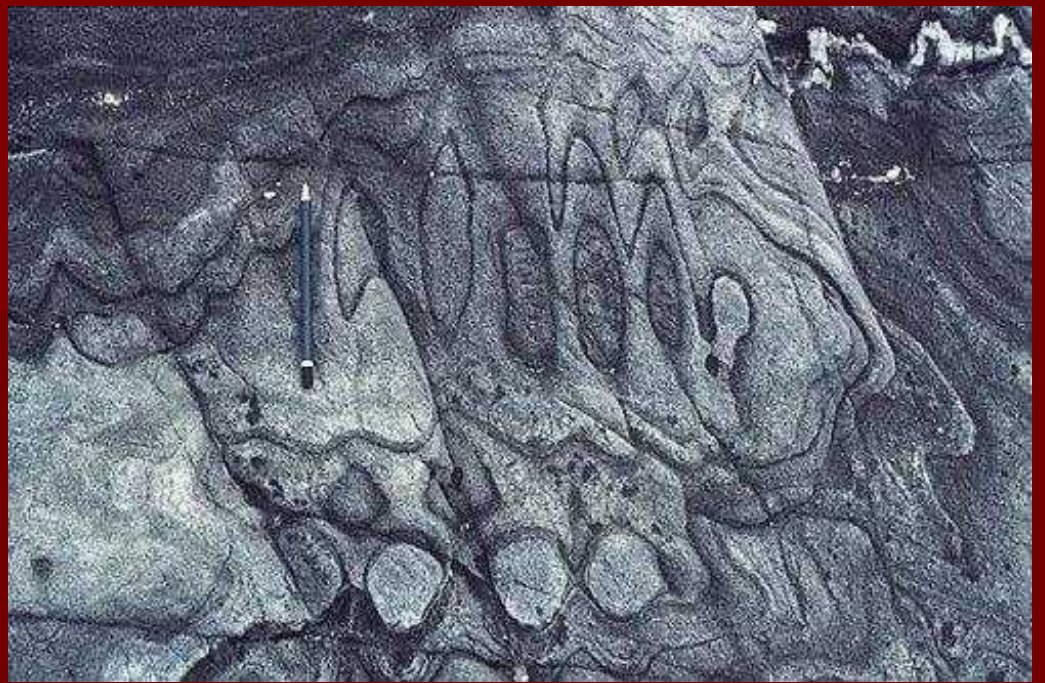
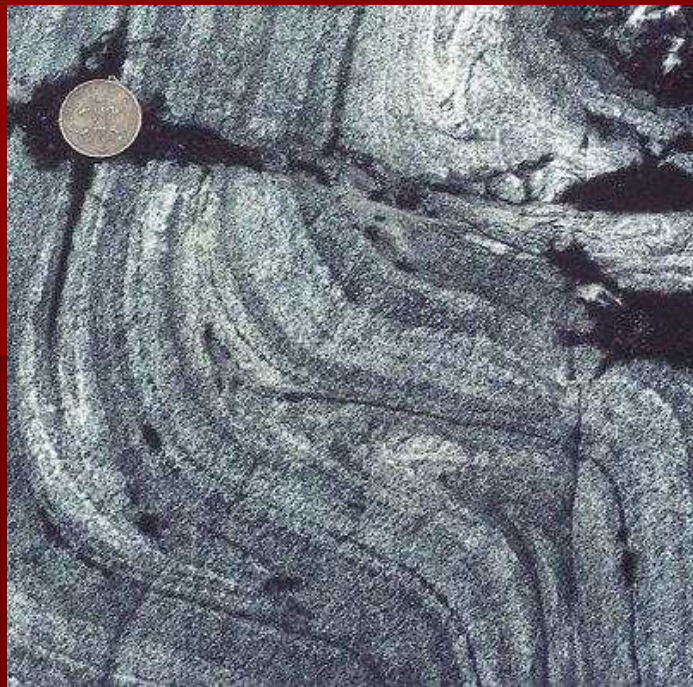
type 2 interference; mushroom



type 3 interference

fold interference patterns





more interference patterns



# Metamorphic Textures

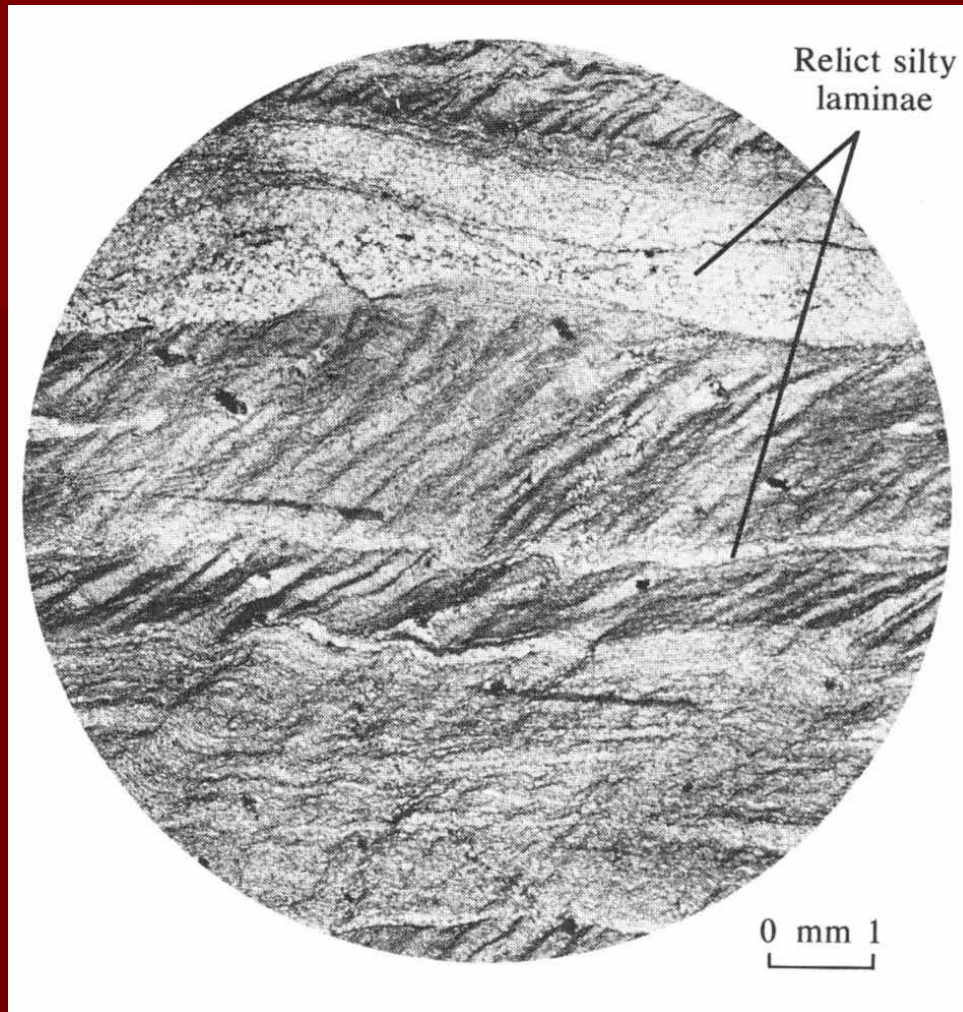
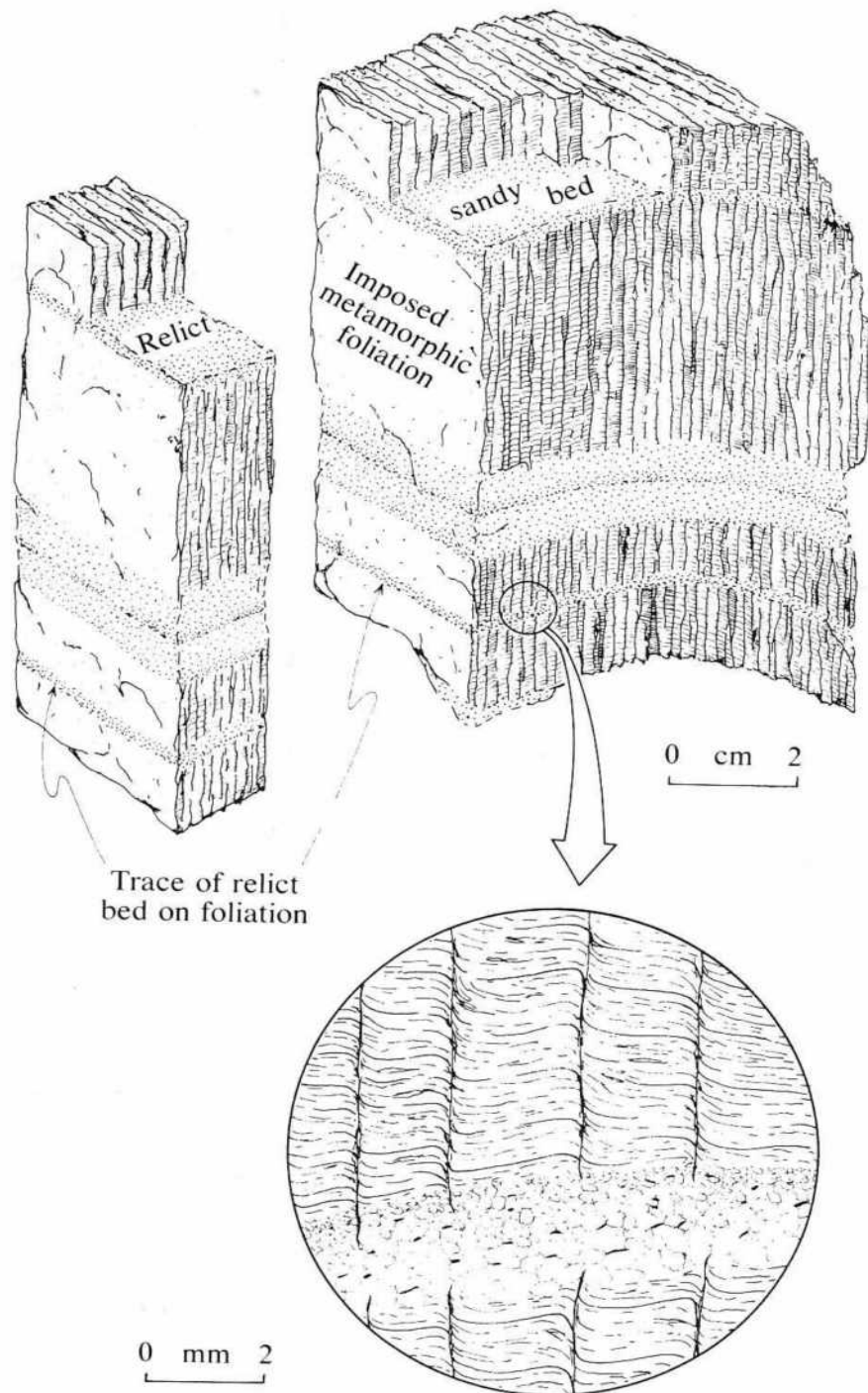
## Textures of Regional Metamorphism

- **Dynamothermal** (crystallization under dynamic conditions)
- **Orogeny**- long-term mountain-building
  - May comprise several **Tectonic Events**
    - May have several **Deformational Phases**
- May have an accompanying **Metamorphic Cycles** with one or more **Reaction Events**

# Textures of Regional Metamorphism

- **Tectonite**- a deformed rock with a texture that records the deformation
- **Fabric**- the complete spatial and geometric configuration of textural elements
  - **Foliation**- planar textural element
  - **Lineation**- linear textural element





**Development of an axial-planar cleavage in folded metasediments. Circular images are microscopic views showing that the axial-planar cleavage is a crenulation cleavage, and is developed preferentially in the micaceous layers.**

# A Morphological Classification of Cleavage and Schistosity (at the thin-section scale)

are microlithons present?

no

yes

(if yes)

can crenulations be recognized  
in the microlithons?

no

yes

**continuous foliation**

if fine grained:  
**continuous  
cleavage**  
or  
**slaty cleavage**



if grains are  
visible to the  
unaided eye:  
**continuous  
schistosity**



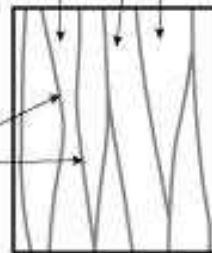
**spaced foliation**

if fine grained:  
**spaced cleavage**

if coarse grained:  
**spaced schistosity**

microlithons

cleavage  
domains



**disjunctive  
foliation /  
cleavage**

**crenulation  
cleavage**

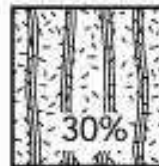
volume % of cleavage domains



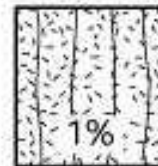
100%



70%



30%



1%

continuous

zonal

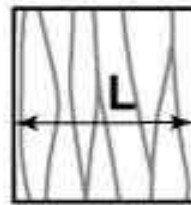
spaced

**A morphological (non-genetic) classification of foliations. After Powell (1979)**



## Other Useful Criteria to Describe Spaced Foliations

### 1. Spacing



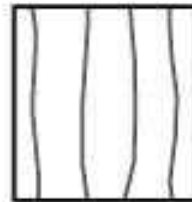
$n=7$

Spacing = length (L) divided by the number of cleavage domains (n) crossed in the length

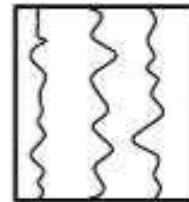
### 2. Shape of cleavage domains



**rough**



**smooth**

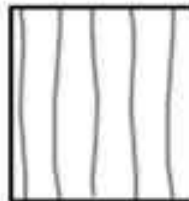


**wiggly**

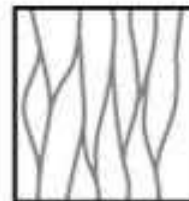


**stylolitic**

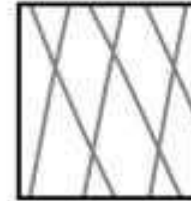
### 3. Spatial relation between cleavage domains



**parallel**

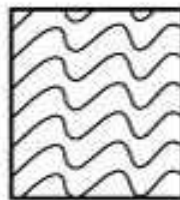


**anastomosing**

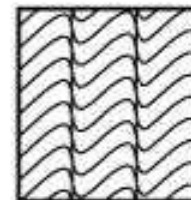
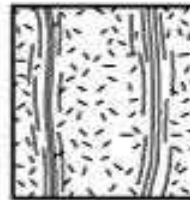


**conjugate**

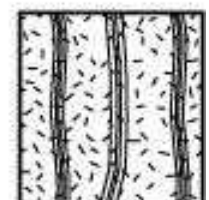
### 4. Transition between cleavage domains and microlithons

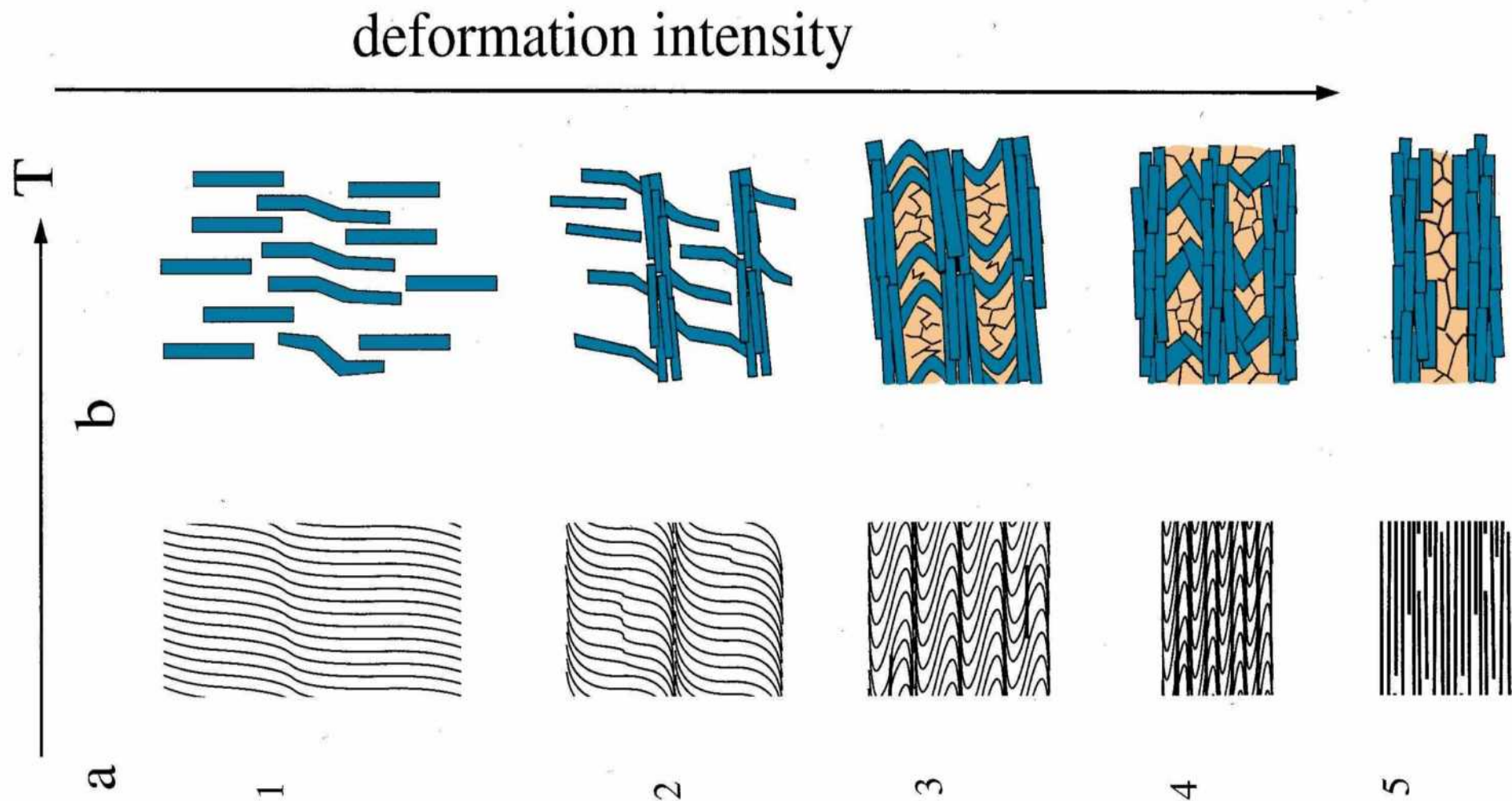


**gradational**



**discrete**

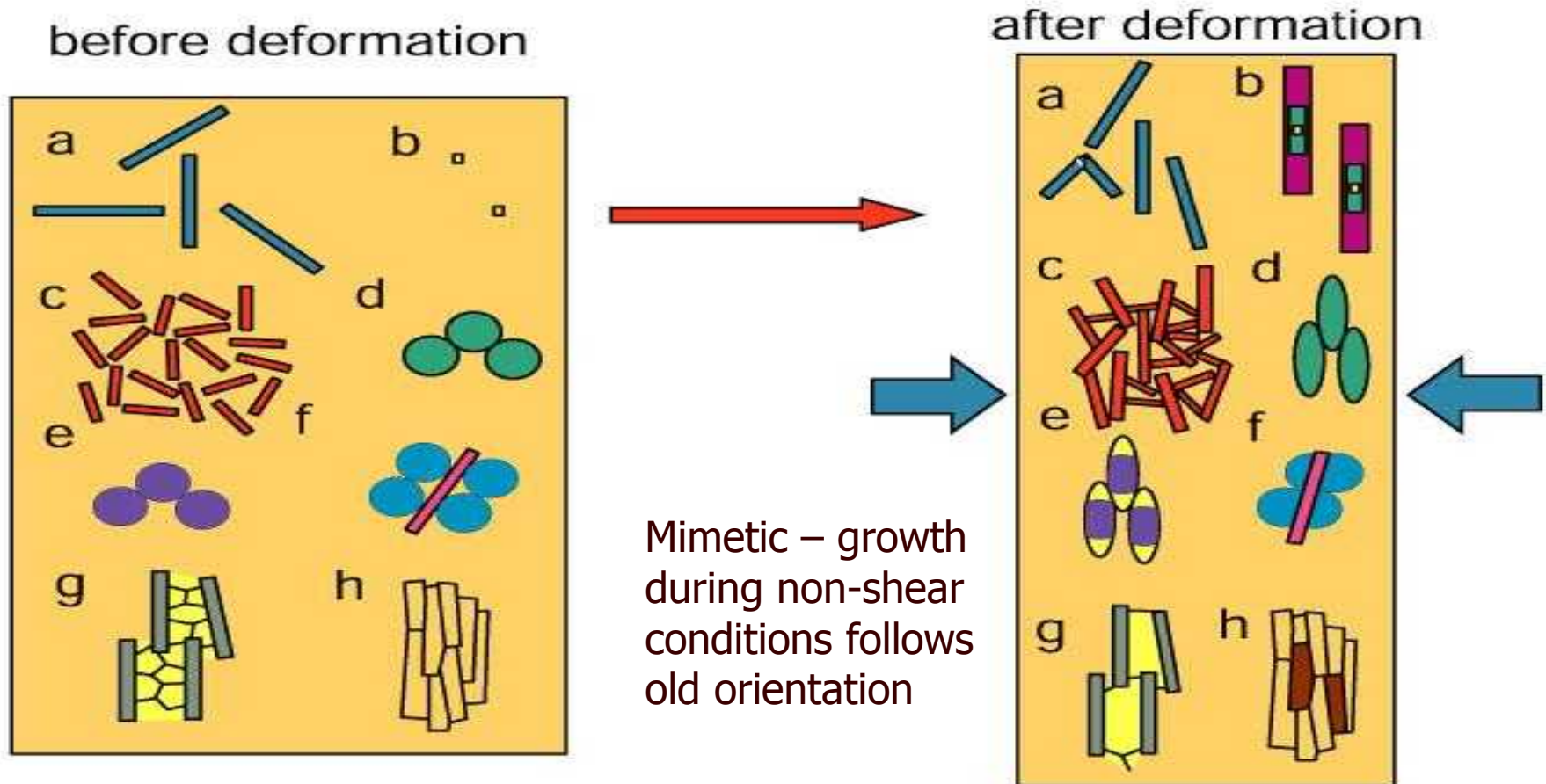




**Development of  $S_2$  micas depends upon T and the intensity of the second deformation**



- Proposed mechanisms for the development of foliations
- a. Mechanical rotation.
  - b. Preferred growth normal to compression.
  - c. Grains with advantageous orientation grow whereas those with poor orientation do not (or dissolve).
  - d. Minerals change shape by ductile deformation.
  - e. Pressure solution.
  - f. A combination of a and e.
  - g. Constrained growth between platy minerals.
  - h. Mimetic growth following an existing foliation.



Development of foliation by simple shear and pure shear (flattening)

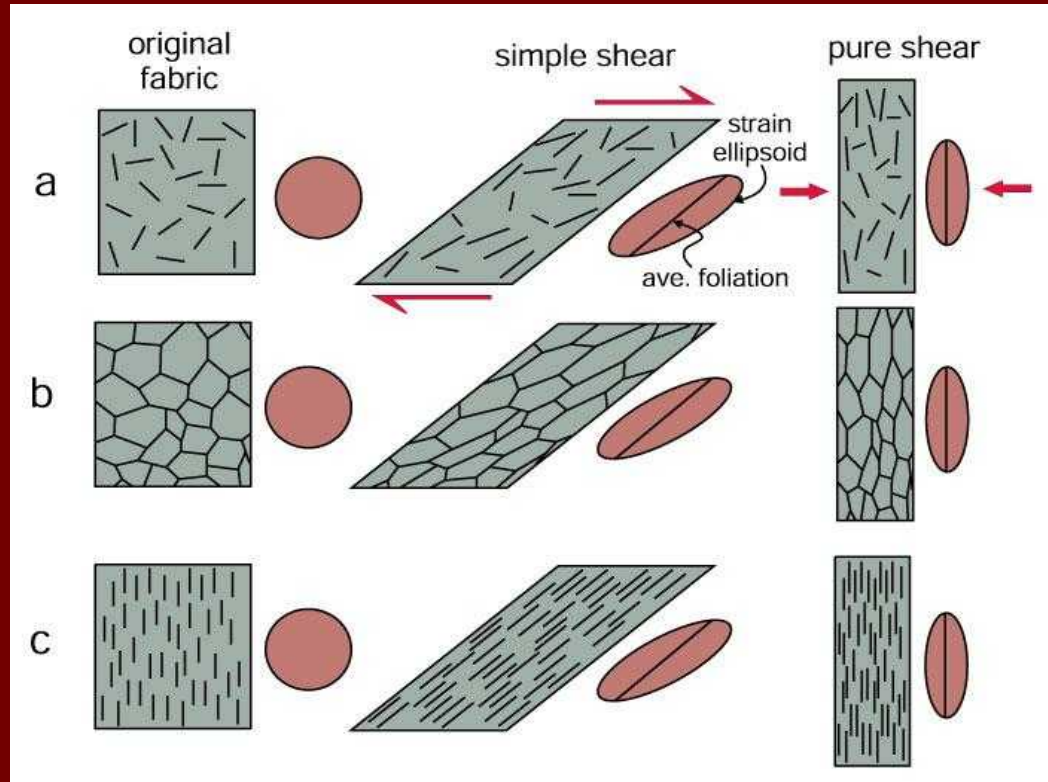
a. Beginning with randomly oriented planar or linear

b. Beginning with equi-dimensional crystals

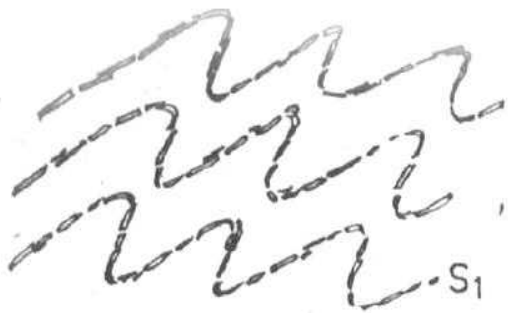
c. Beginning with pre-existing foliation

Shaded figures represent an initial sphere and the resulting strain ellipsoid

**Development of foliation by simple shear and pure shear (flattening).**



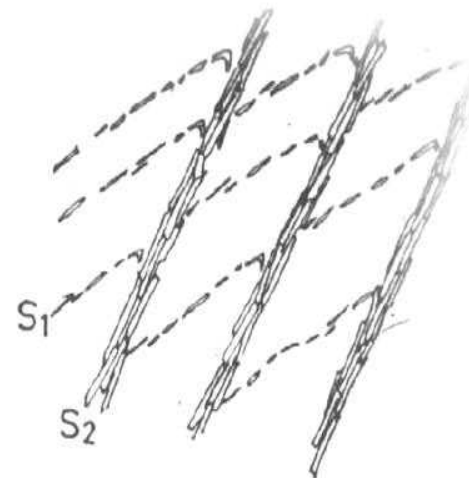




( a )



( b )

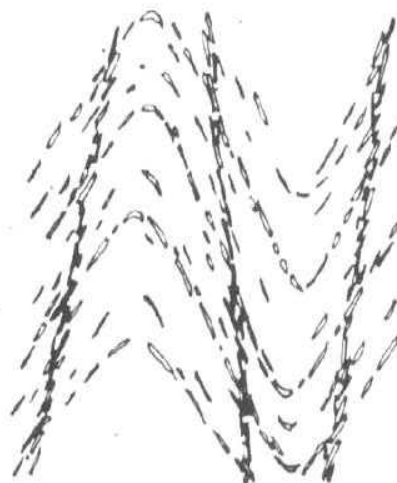


( c )

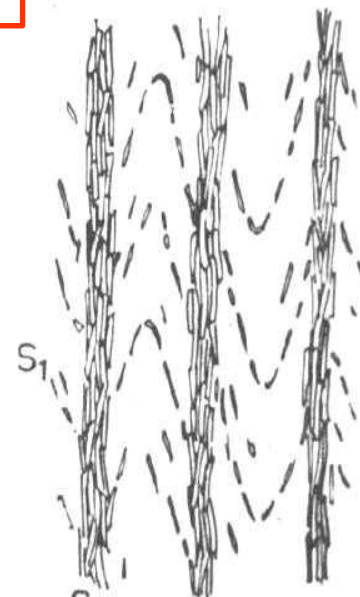
## DEVELOPMENT OF CRENULATION CLEAVAGES



( d )



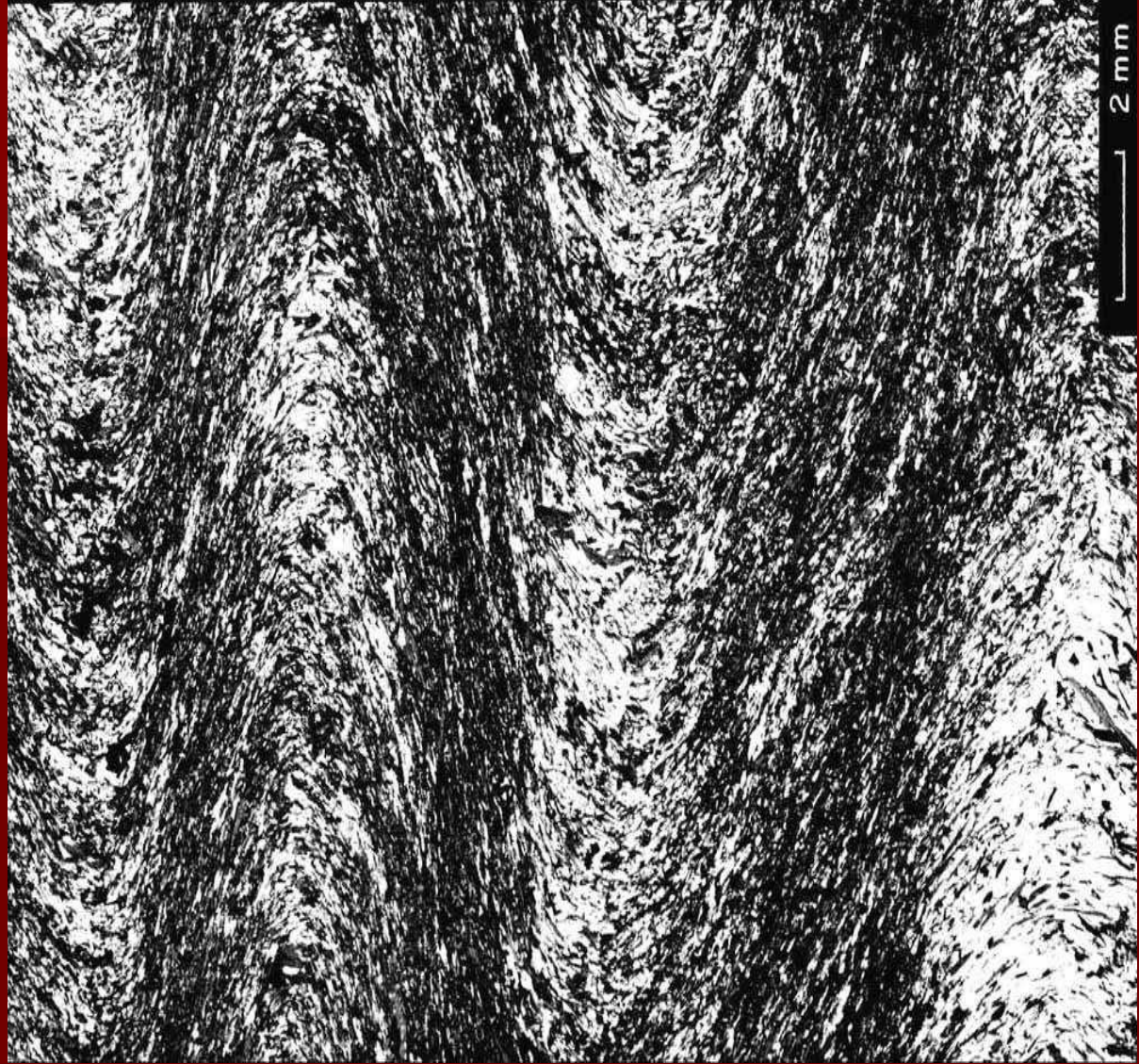
( e )



( f )

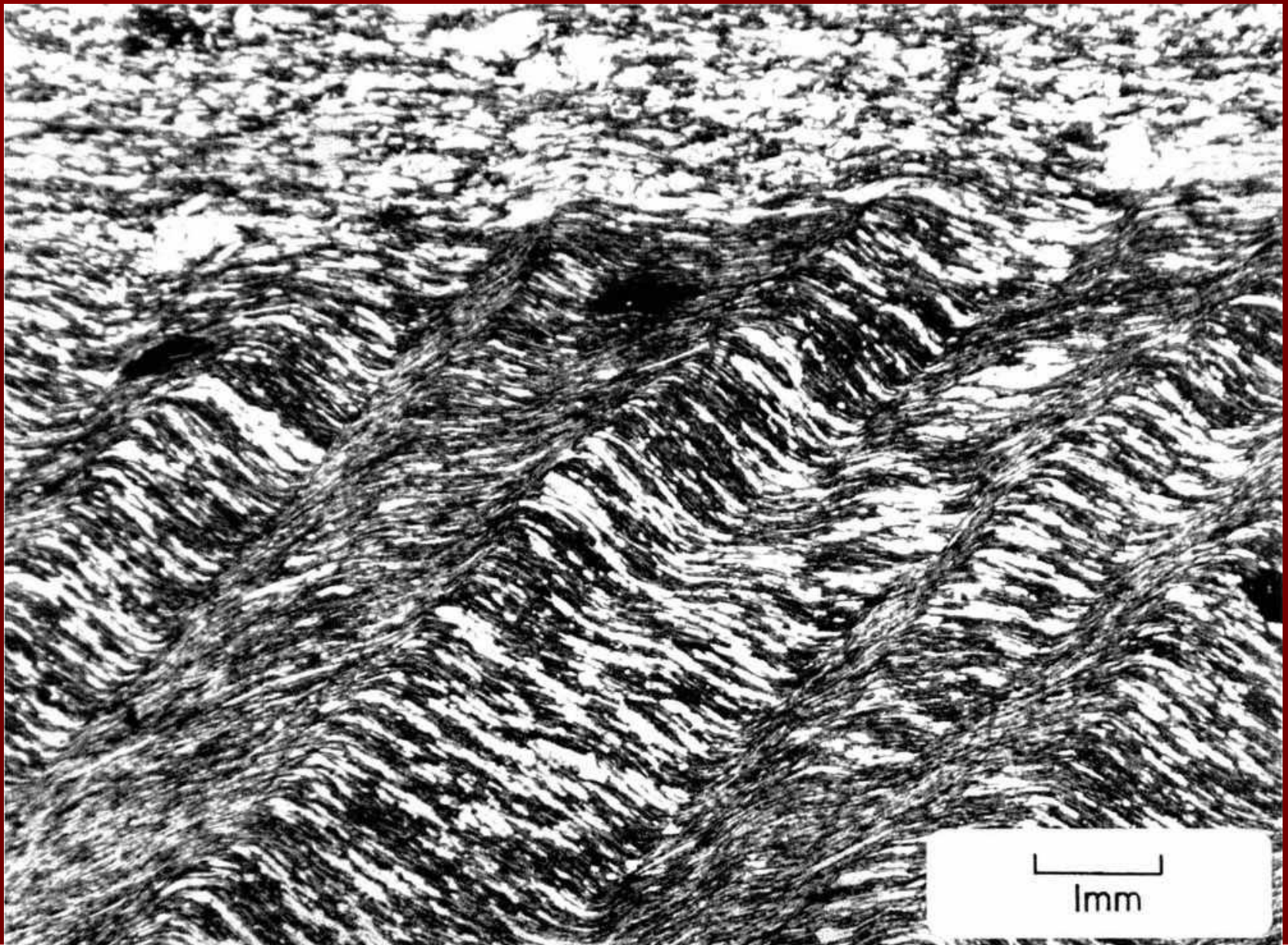
Crenulation: A slaty cleavage or schistosity that becomes microfolded

Quartz grains commonly dissolve by pressure solution from the steep limbs and precipitate in the hinge



**Symmetrical crenulation cleavages** in amphibole-quartz-rich schist. Note concentration of quartz in hinge areas.





**Asymmetric crenulation cleavages in mica-quartz-rich schist. Note horizontal compositional layering (relict bedding) and preferential dissolution of quartz from one limb of the folds.**



# ■ Progressive Syntectonic Metamorphism of a volcanic graywacke.



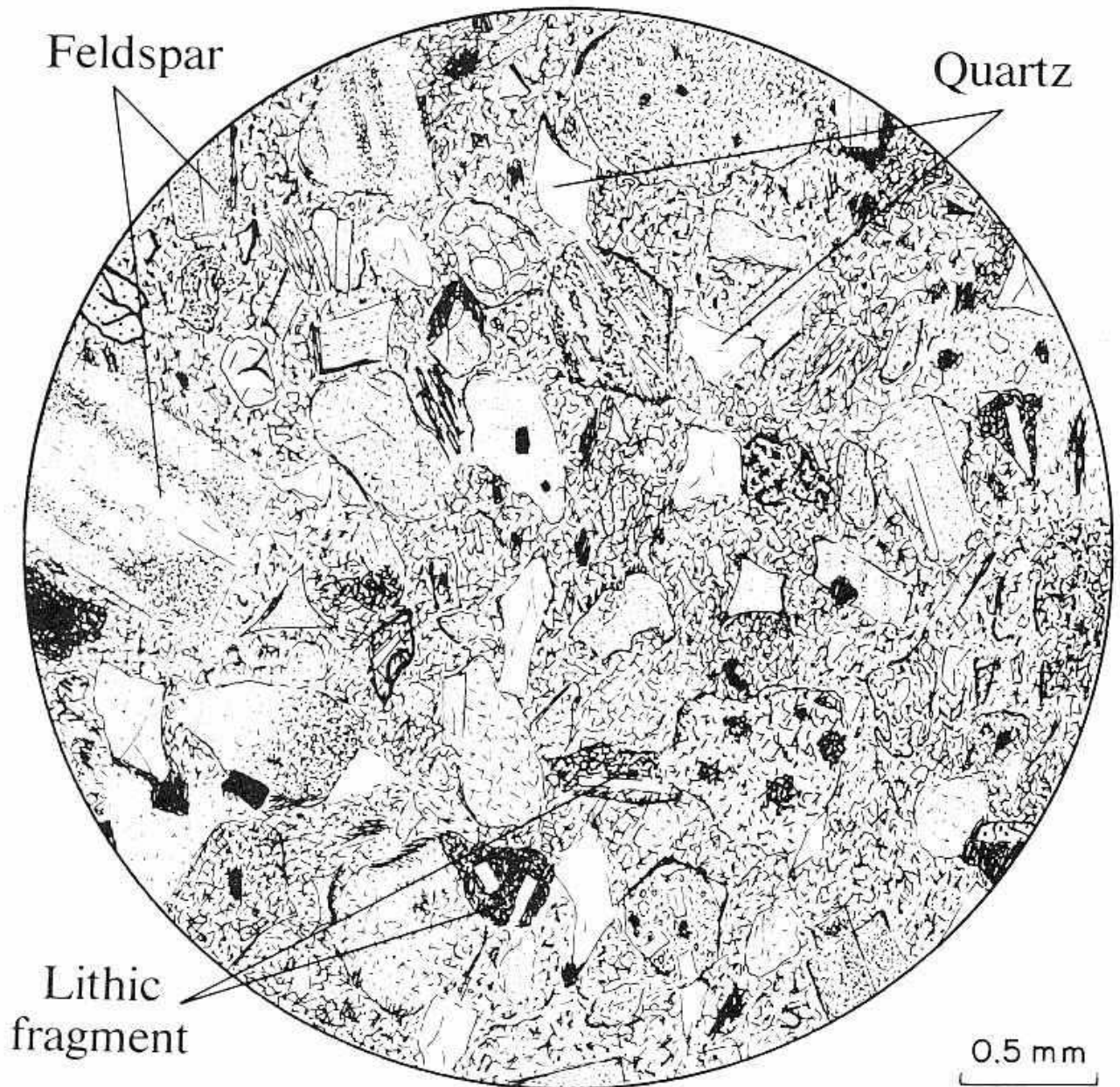


Phase 1: Both plag and Kfs in isotropic matrix

- Feldspars altered to fine sericite and secondary Ca-Al silicates

- Sericite is a fine grained mica, either muscovite or Illite. Sericite is a common alteration mineral of orthoclase or plagioclase feldspars

## Progressive syntectonic metamorphism

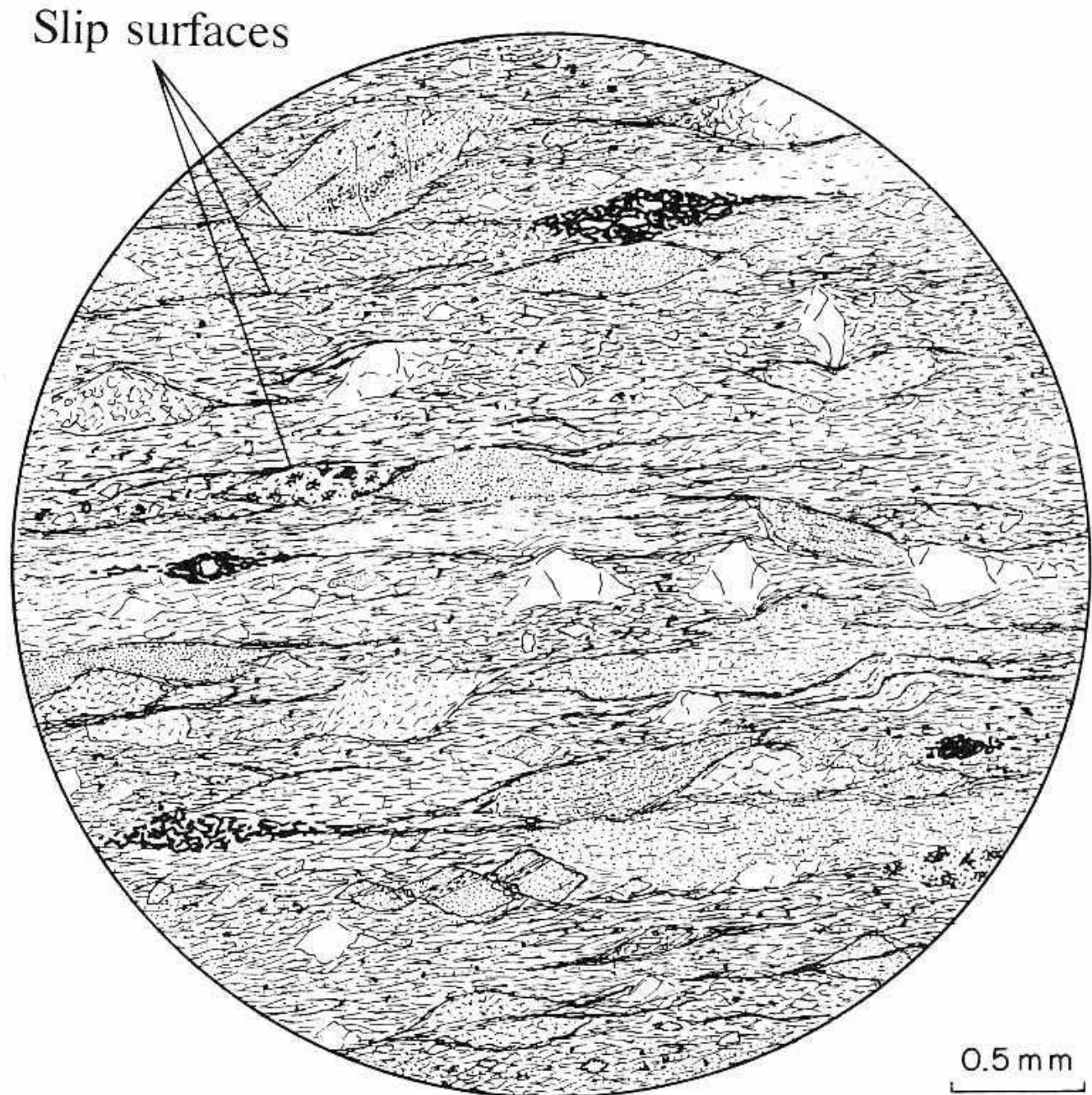




Phase 2: Pervasive foliation develops due mostly to shear

- Grain size reduction
- Porphyroclasts common and rounded
- Matrix recrystallized and new minerals form (Qtz, Ep, Sericite, Ab, Chl)
- Chl & Ser enhance foliation

**Progressive  
syntectonic  
metamorphism**

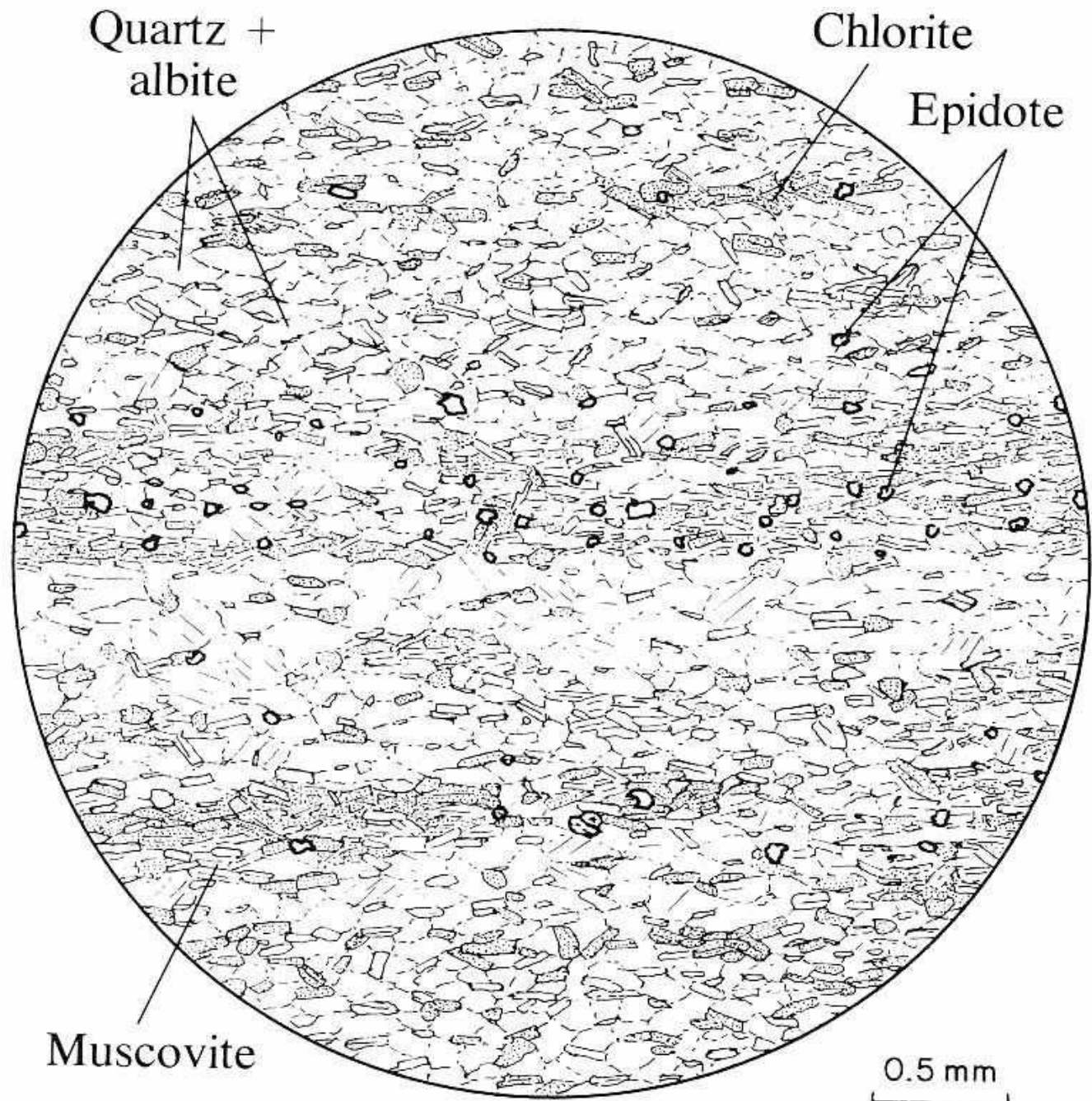




Phase 3: Fine-grained schist with larger crystals- no relict textures

- Good muscovite and biotite define schistosity
- Some metamorphic differentiation to layering
- Qtz and Ab are polygonal mosaic in mica-free layers

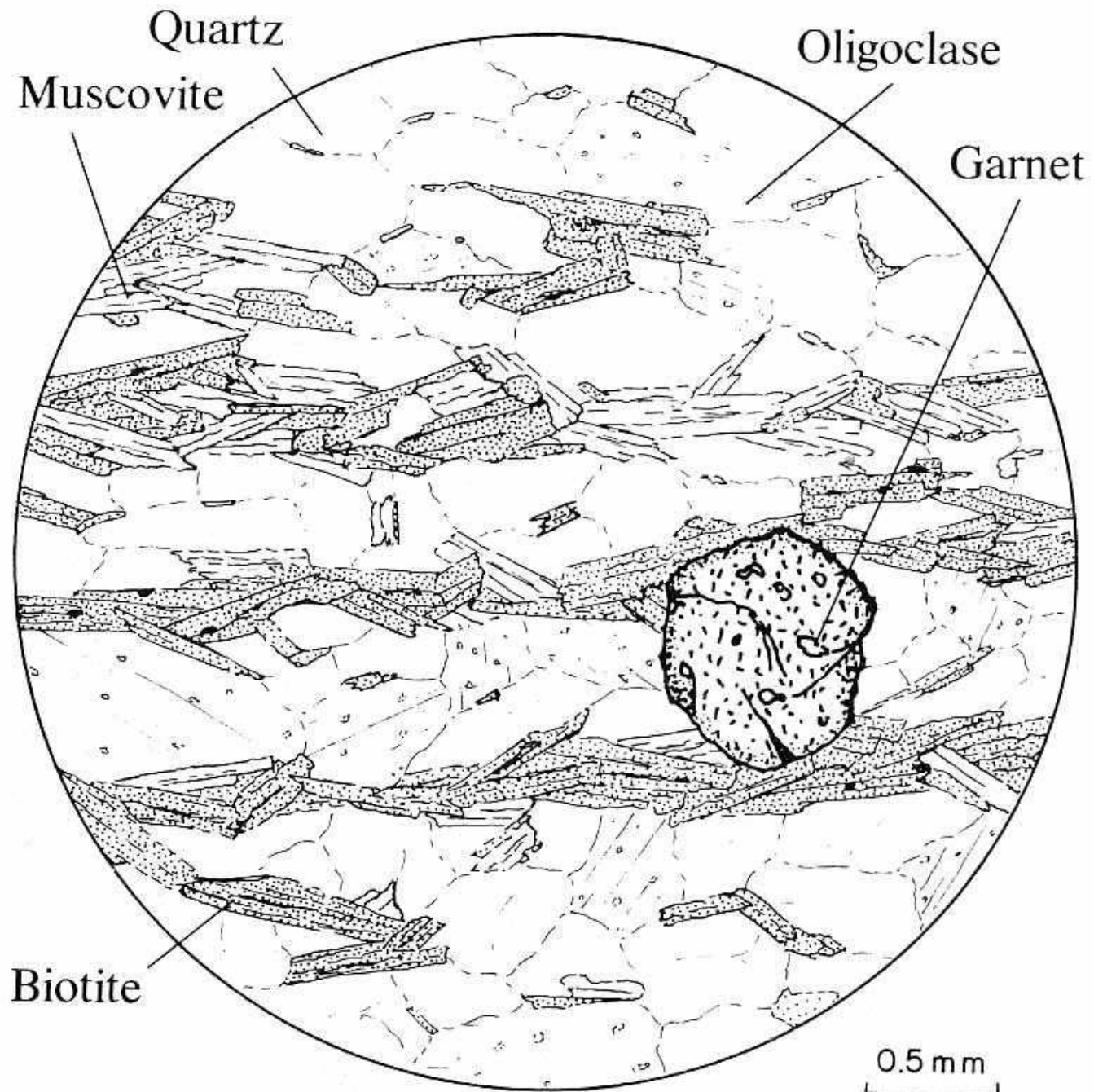
**Progressive  
syntectonic  
metamorphism**



Phase 4: Good schist with coarser grains

- More enhanced segregation into layers
- Plag no longer Ab- accepts more Ca at higher T
- Garnet is a new isograd mineral

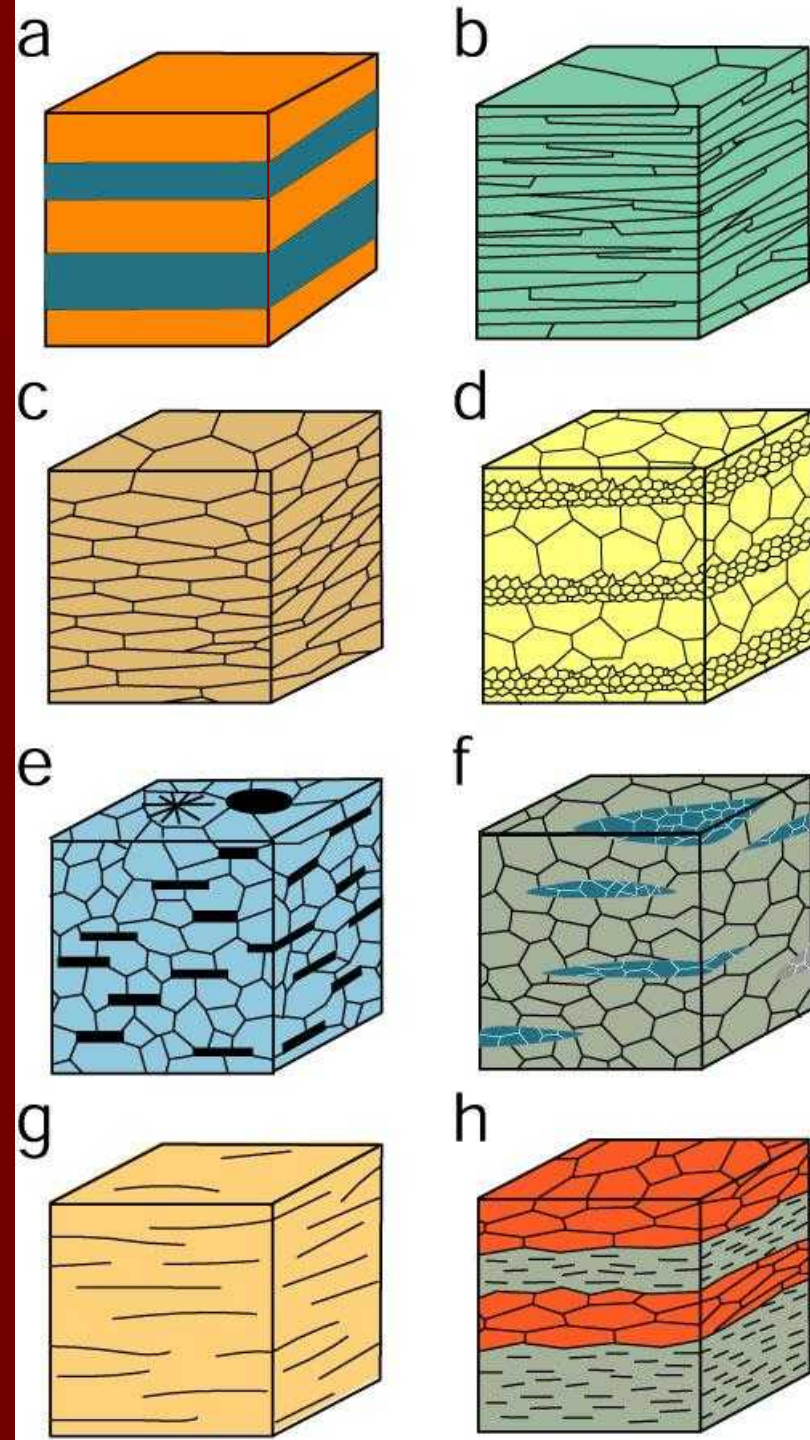
**Progressive  
syntectonic  
metamorphism of  
a volcanic  
graywacke,**





# Types of foliations

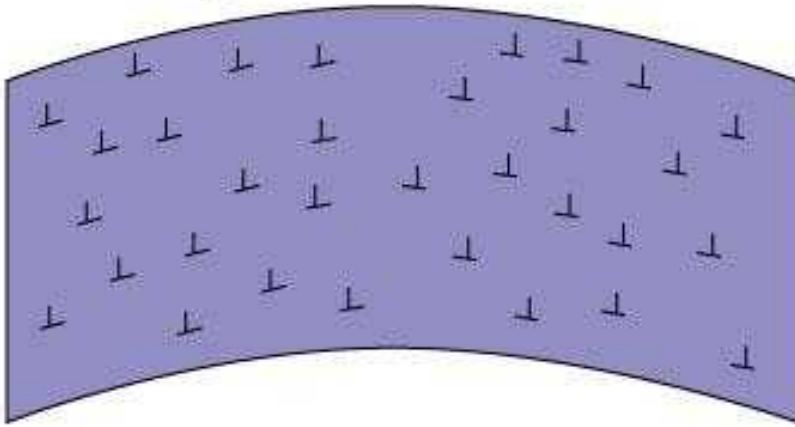
- a.** Compositional layering
- b.** Preferred orientation of platy minerals
- c.** Shape of deformed grains
- d.** Grain size variation
- e.** Preferred orientation of platy minerals in a matrix without preferred orientation others
- f.** Preferred orientation of lenticular mineral aggregates
- g.** Preferred orientation of fractures
- h.** Combinations of the above



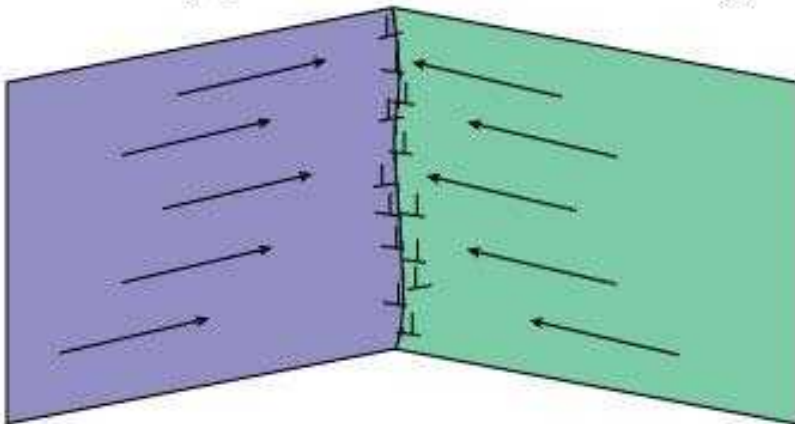
# Textural Effects of Differential Stress

Stressed grains store strain energy released during metamorphism  
clearing defects and dislocations (polygonization)

a. strained grain with undulose extinction



b. recovery produces two strain-free subgrains



Undulose extinction



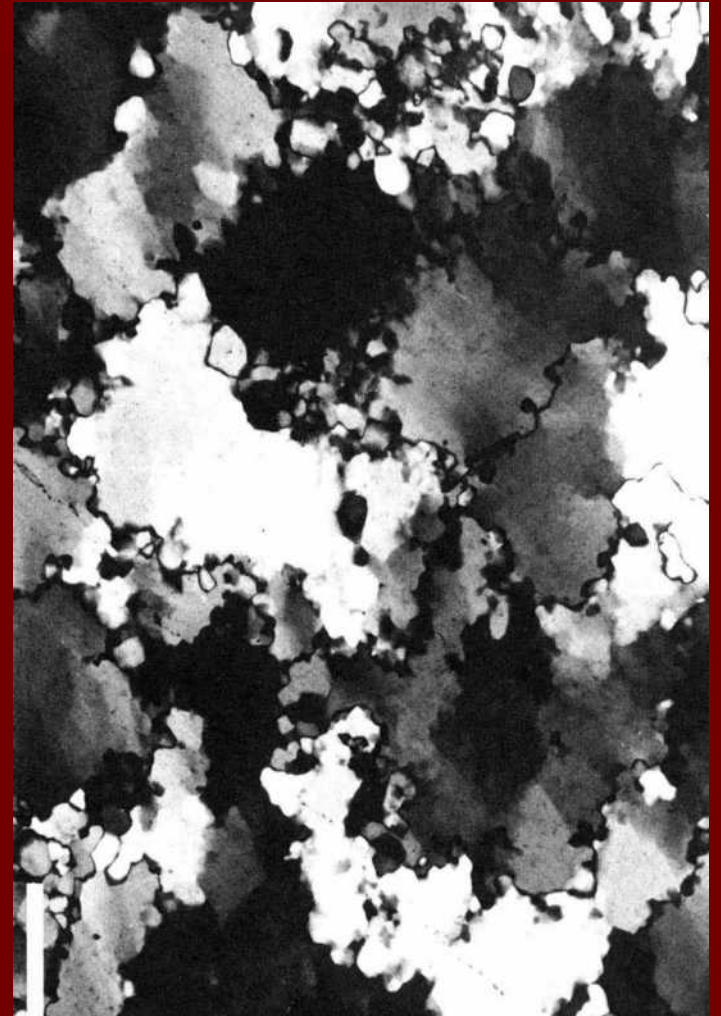
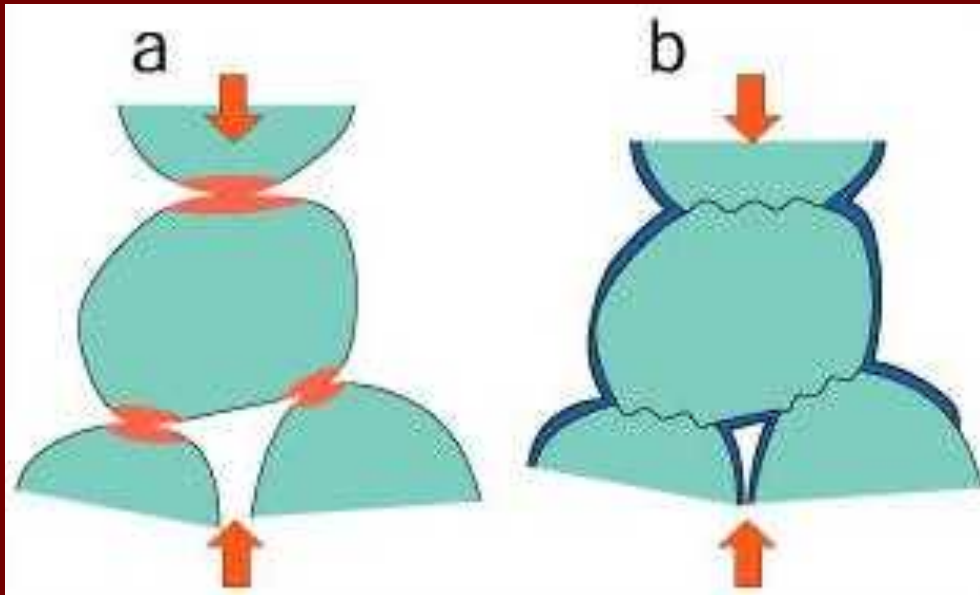
Subgrain development





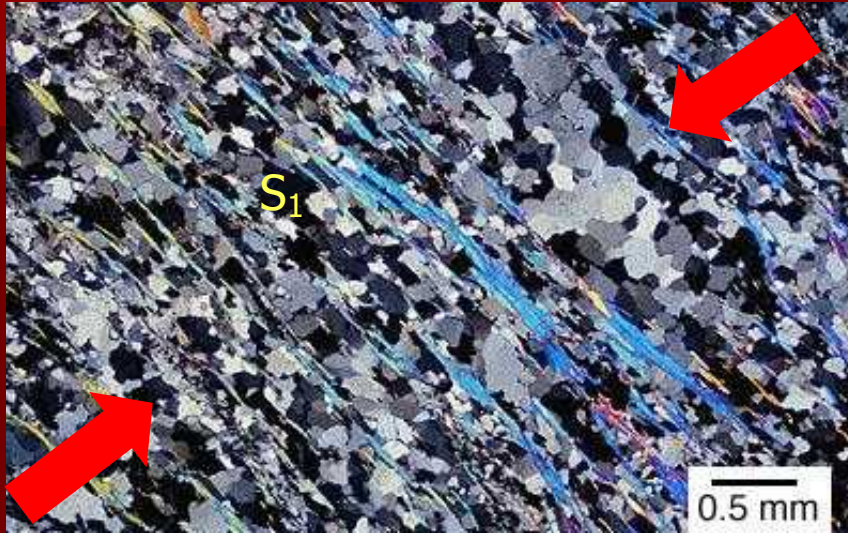
# Pressure Solution

Differential stress produces serrated grain boundaries



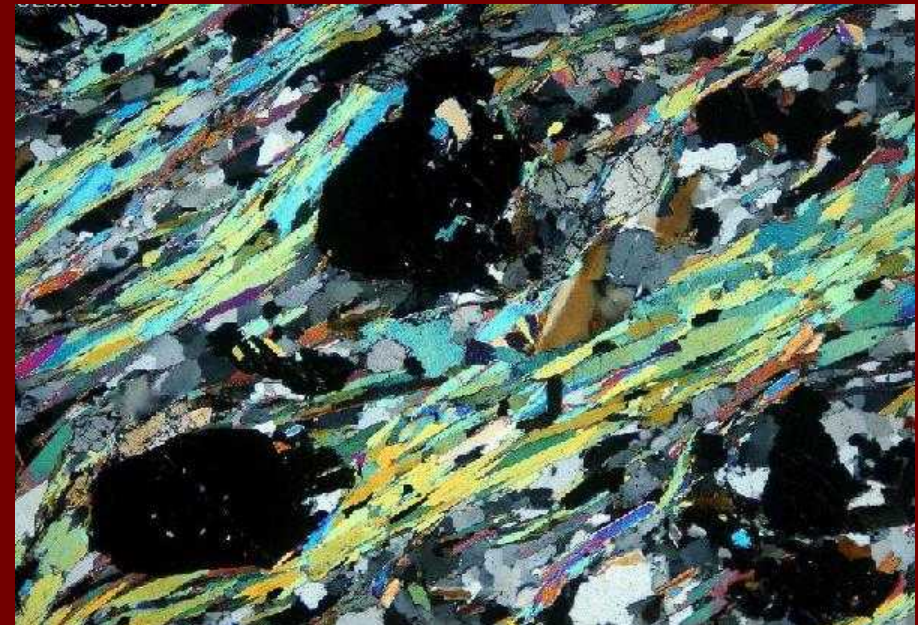
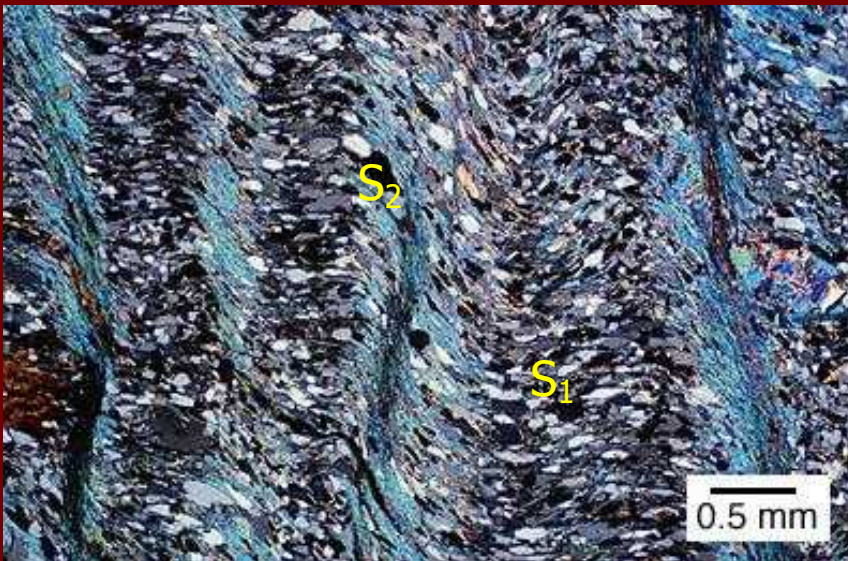


# Foliation Development



$\sigma_1$

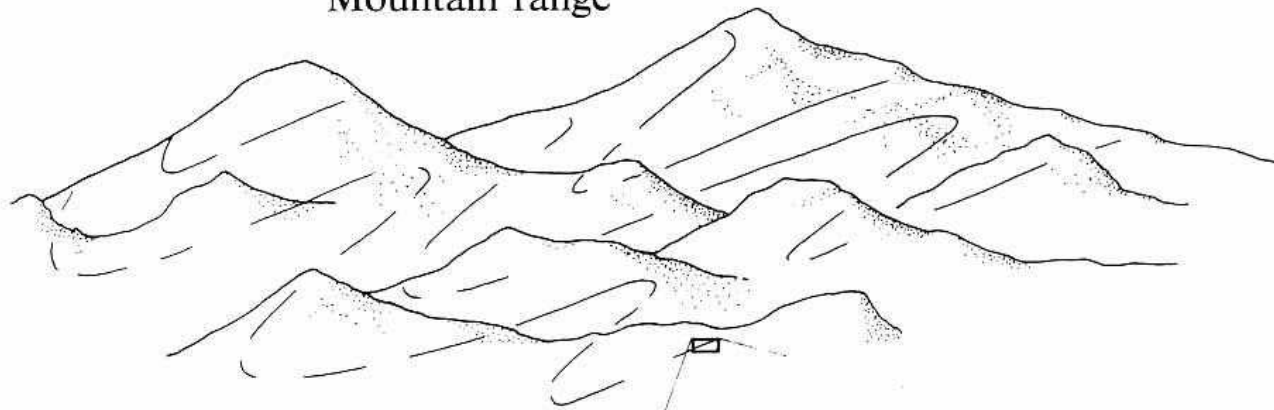
Schistosity  $S_1$  flattened around garnet porphyroblasts



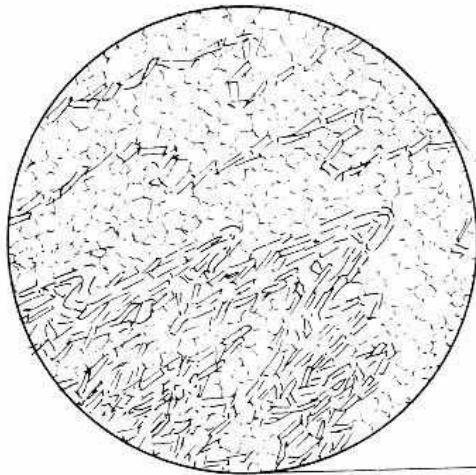
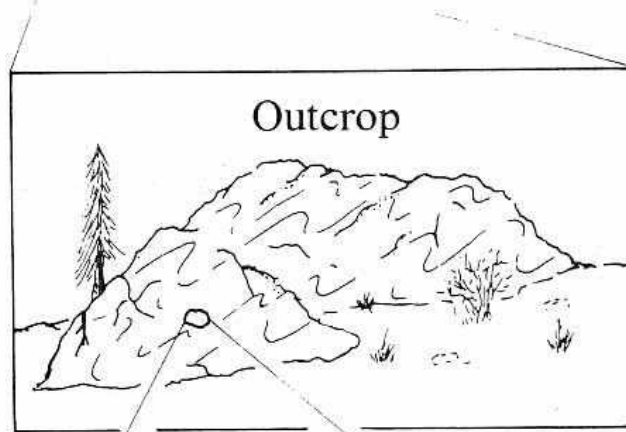
Crenulation cleavage  $S_2$   
developed by folding  $S_1$



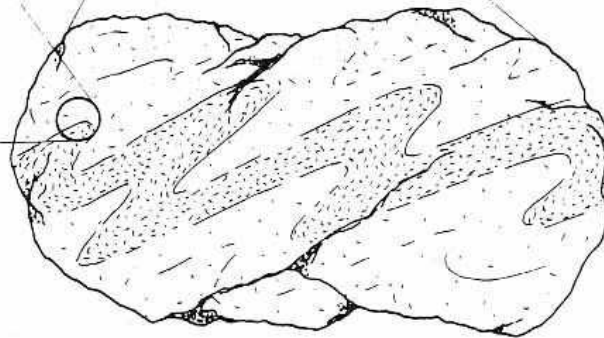
Mountain range



Outcrop

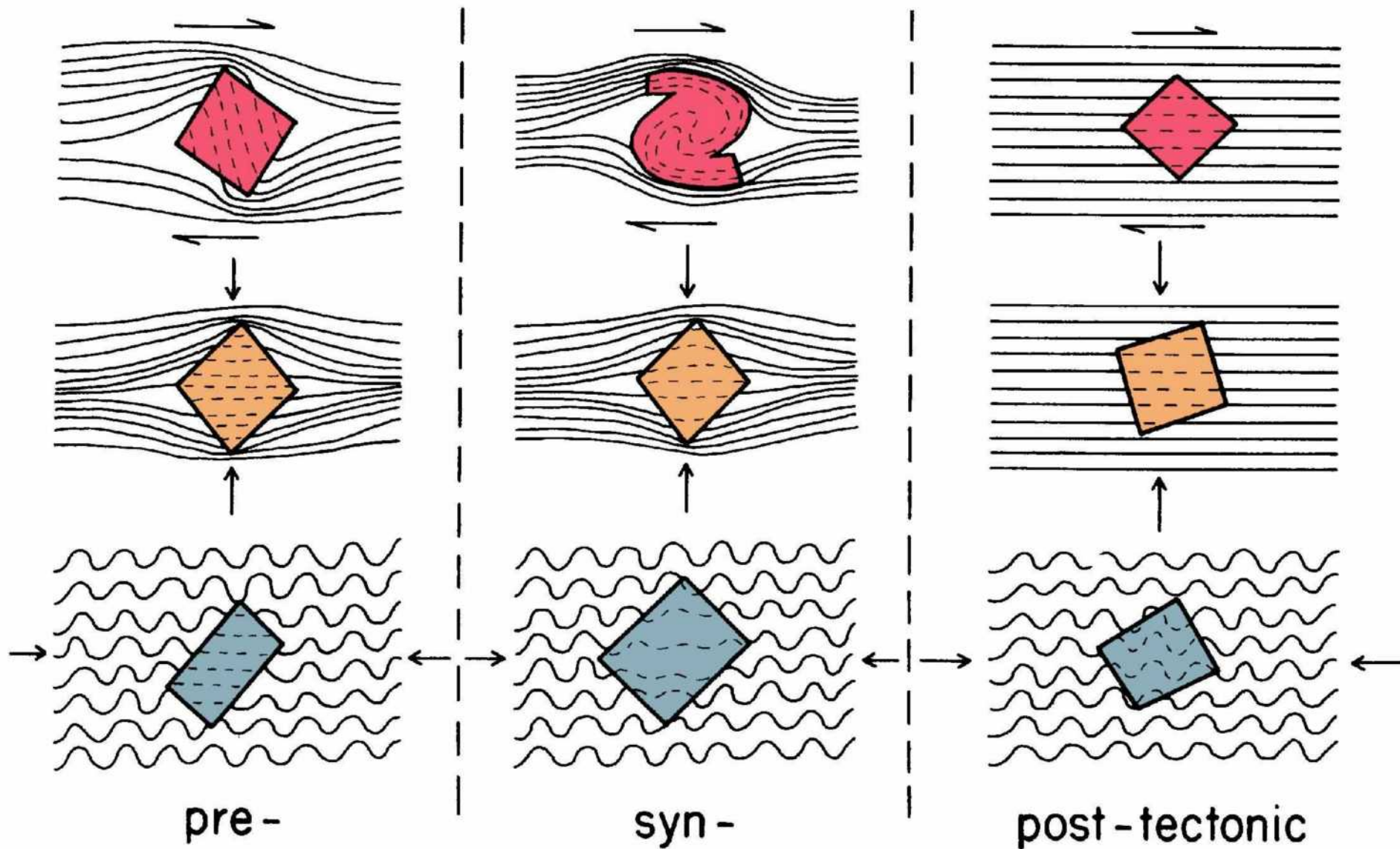


Thin section



Hand sample

**Diagram showing that structural and fabric elements are generally consistent in style and orientation at all scales.**

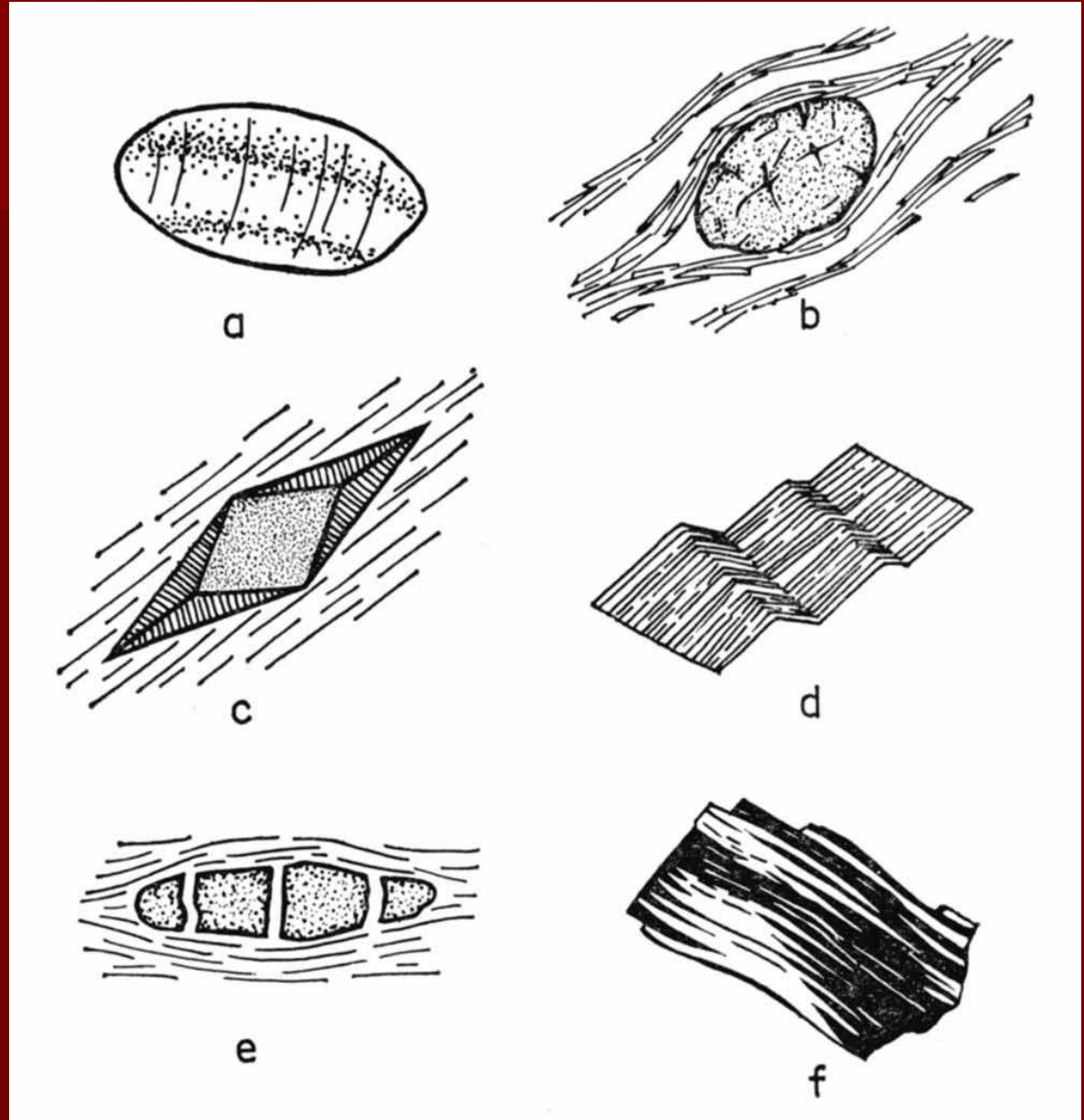


**$S_1$  characteristics of clearly pre-, syn-, and post-kinematic crystals as proposed by Zwart (1962). a. Progressively flattened  $S_1$  from core to rim. b. Progressively more intense folding of  $S_1$  from core to rim. c. Spiraled  $S_1$  due to rotation of the matrix or the porphyroblast during growth. After Zwart (1962) *Geol. Rundschau*, 52, 38-65.**

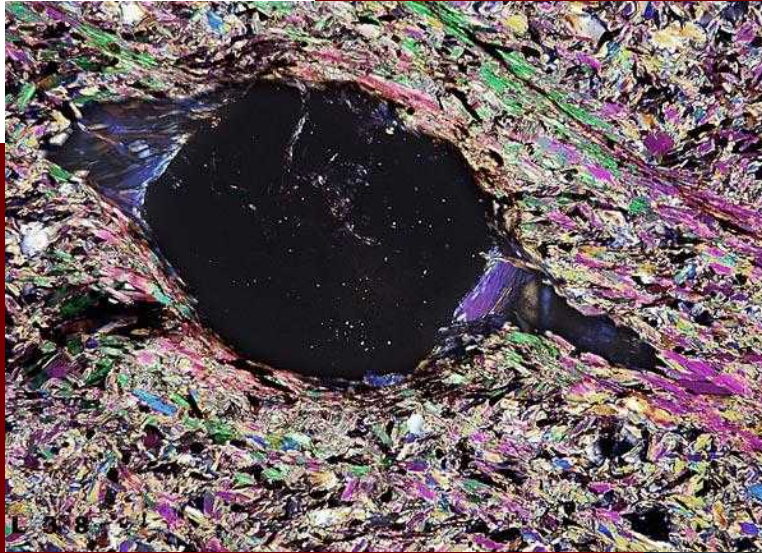
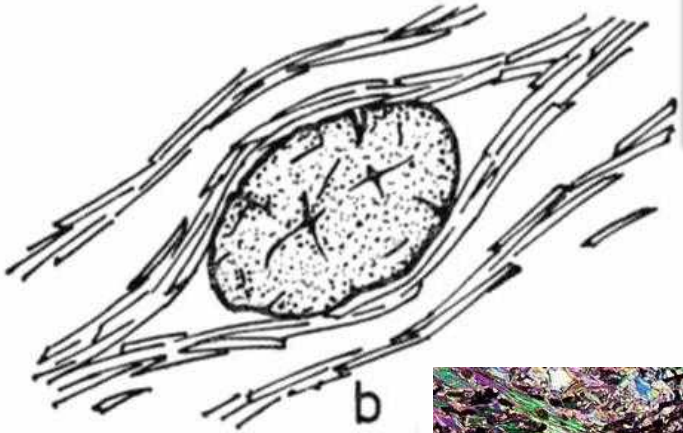


# Pre-kinematic crystals

- a. Bent crystal with undulose extinction
- b. Foliation wrapped around a porphyroblast
- c. Pressure shadow or fringe
- d. Kink bands or folds
- e. Microboudinage
- f. Deformation twins



# Pressure Shadows



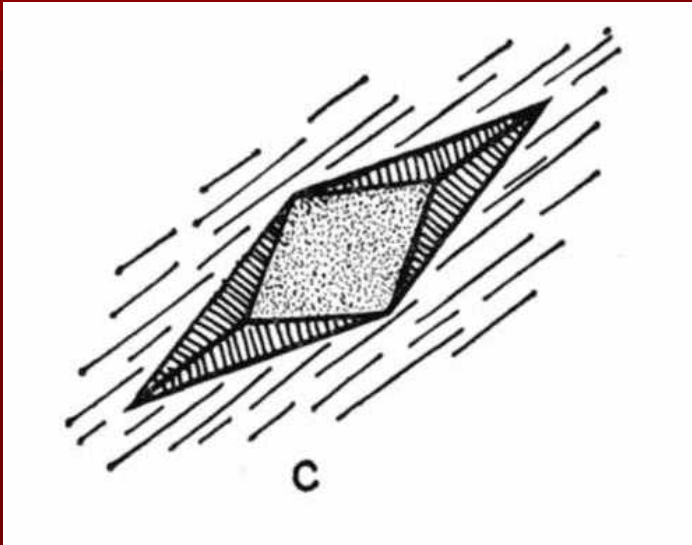
Chlorite in shadow  
around garnet

Quartz in shadow  
around staurolite





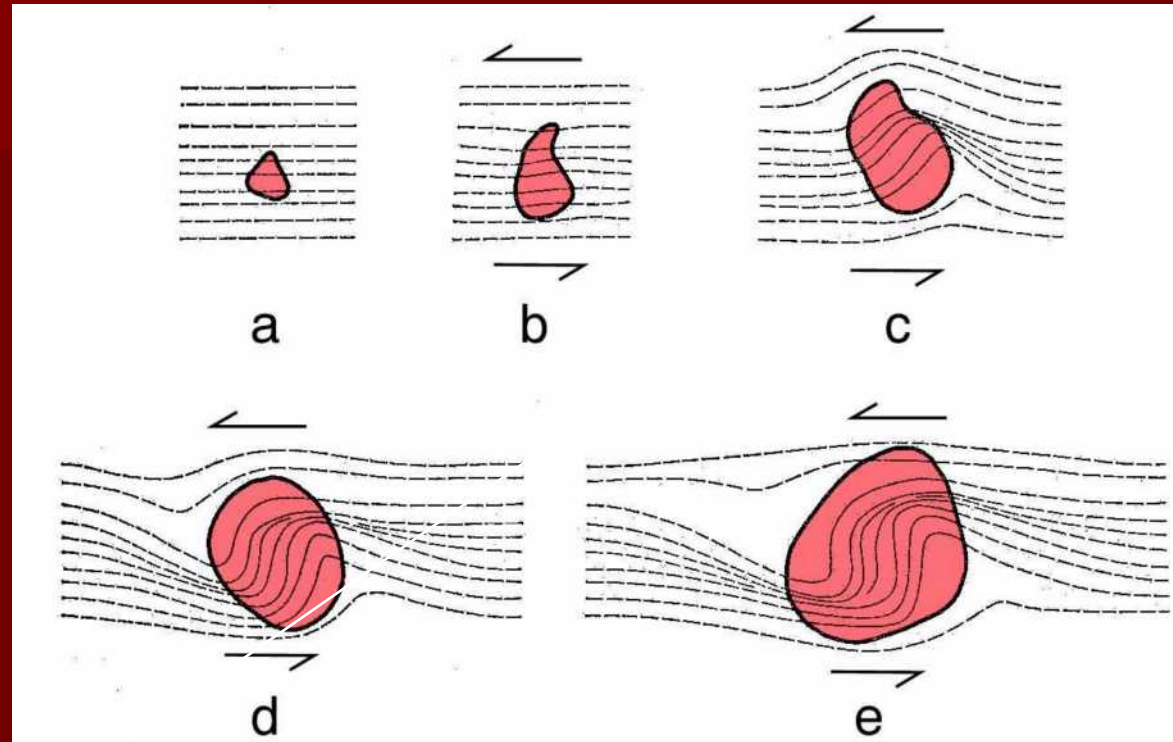
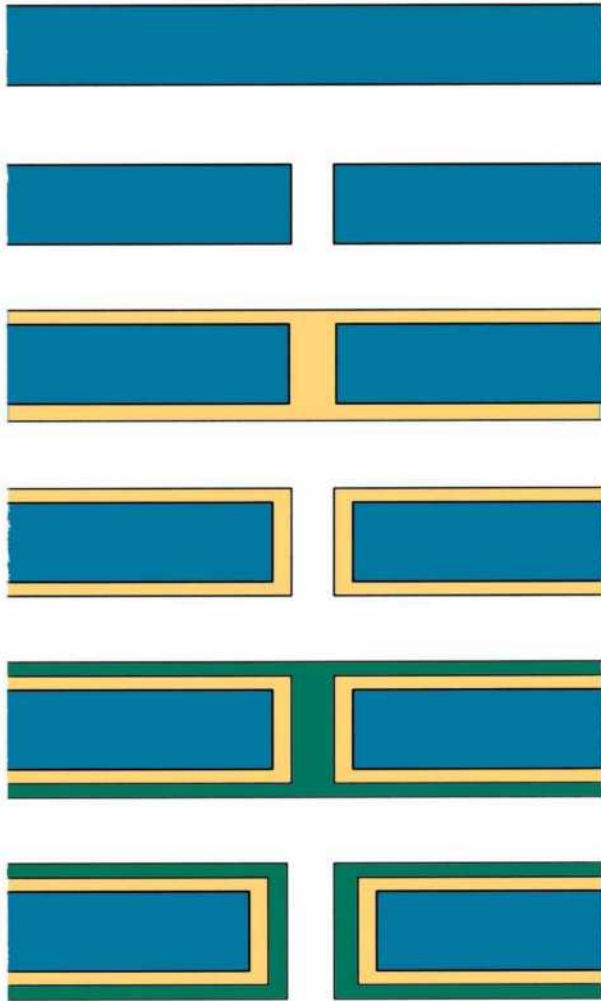
# Pressure Fringes



# Syn-kinematic crystals

Paracrystalline microboudinage

Spiral Porphyroblast

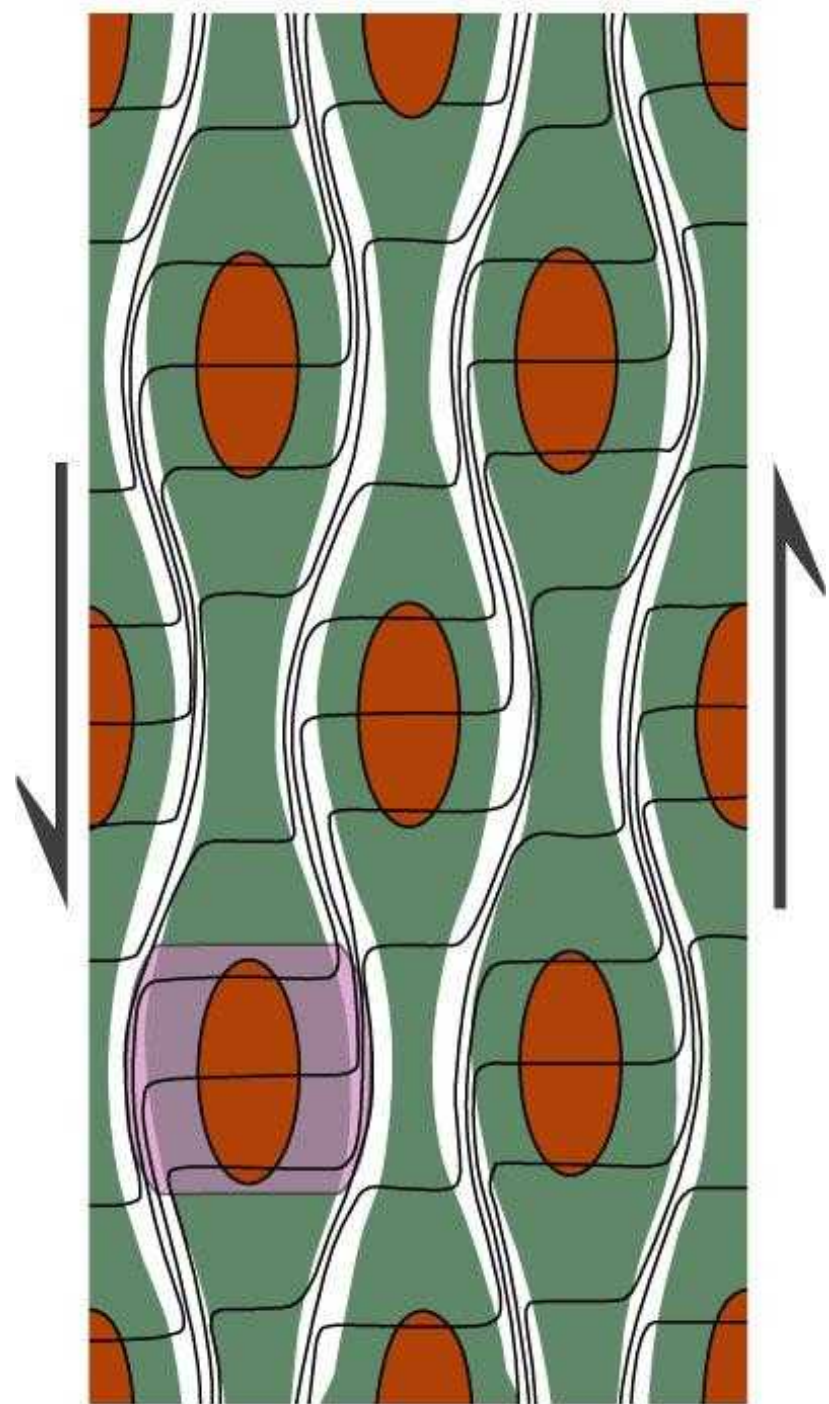


Traditional interpretation of spiral  $S_1$  train in which a porphyroblast is rotated by shear as it grows.

**Syn-crystallization micro-boudinage. Syn-kinematic crystal growth can be demonstrated by the color zoning that grows and progressively fills the gap between the separating fragments.**



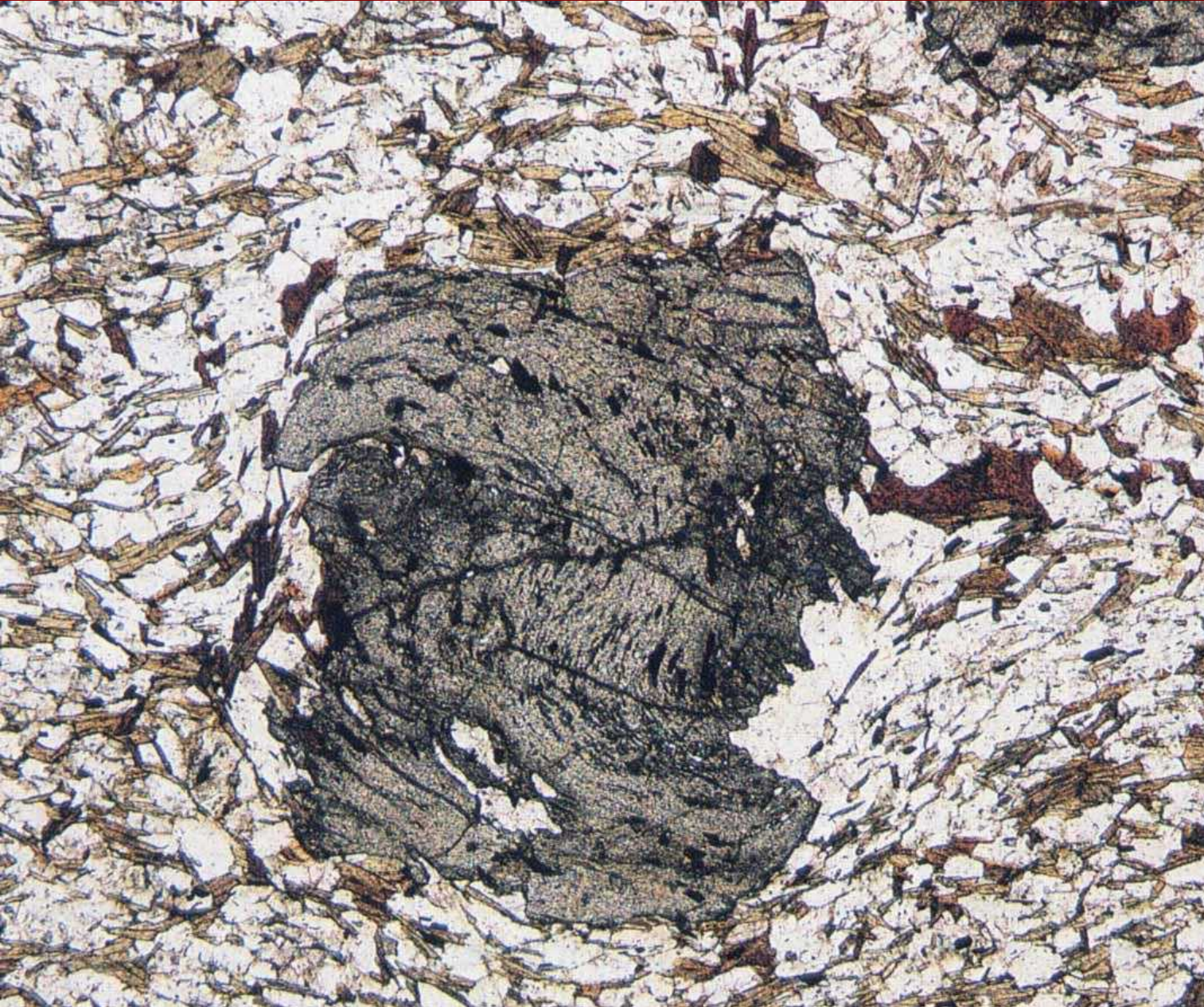
# Syn-kinematic crystals



**Figure 23.40.** Non-uniform distribution of shear strain as proposed by Bell *et al.* (1986) *J. Metam. Geol.*, 4, 37-67. Blank areas represent high shear strain and colored areas are low-strain. Lines represent initially horizontal inert markers ( $S_1$ ). Note example of porphyroblast growing preferentially in low-strain regions.



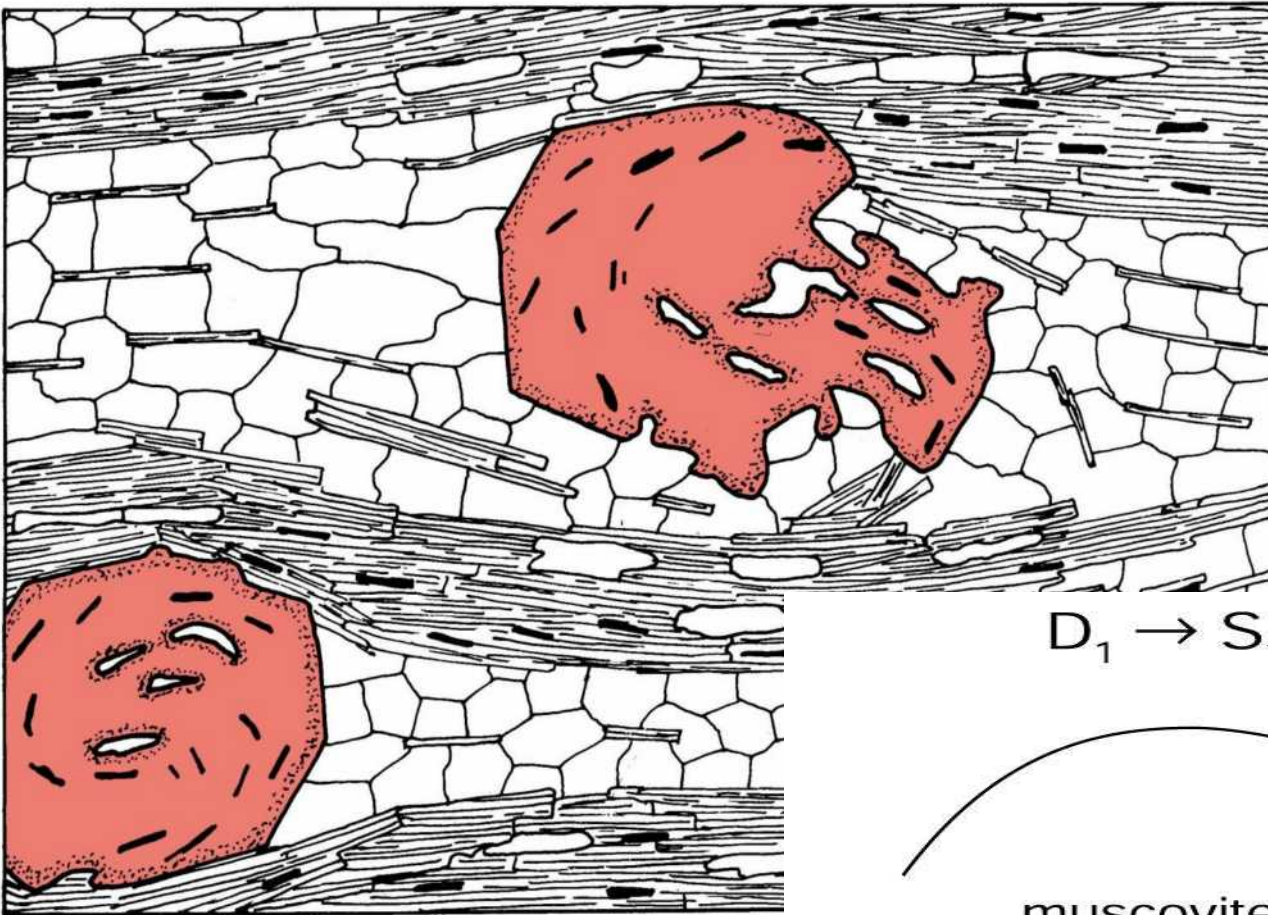
# Syn-kinematic crystals



**Spiral  
 $S_1$   
trails  
in  
garnet**



# Syn-tectonic Texture



$D_1 \rightarrow S_1$

rotation of  $S_1$   
↔

muscovite

biotite

quartz

garnet

Time →



## Snowball Garnet with Pressure Shadows





# Syn-kinematic crystals

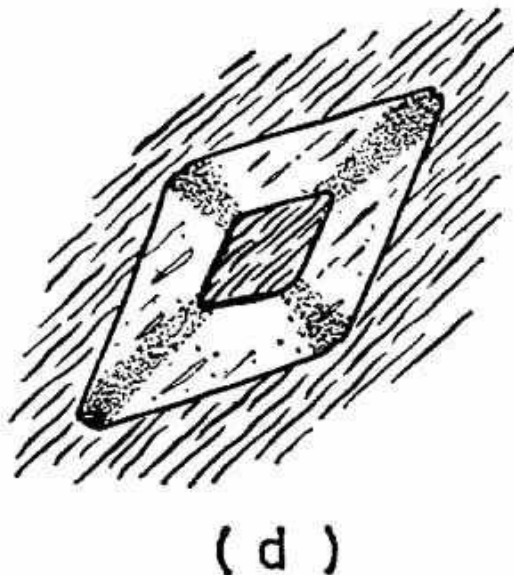
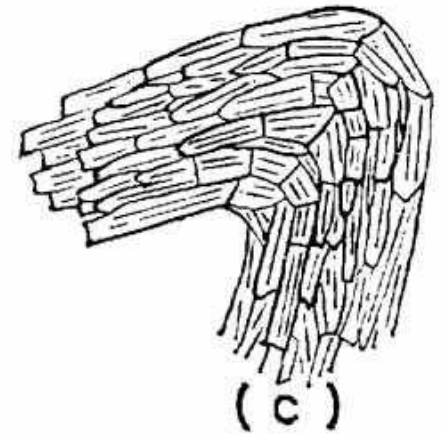
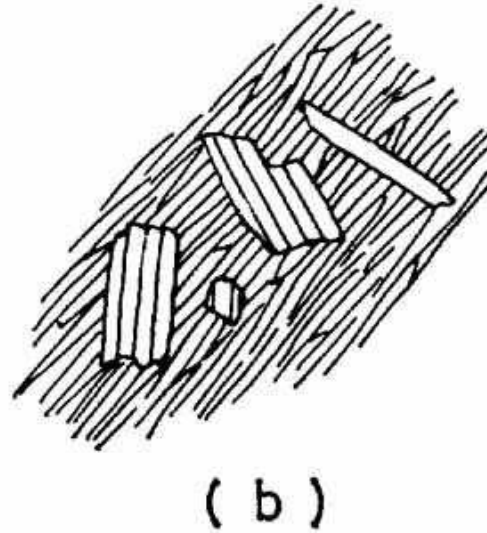


**"Snowball  
garnet"  
with  
highly  
rotated  
spiral  $S_1$ .**

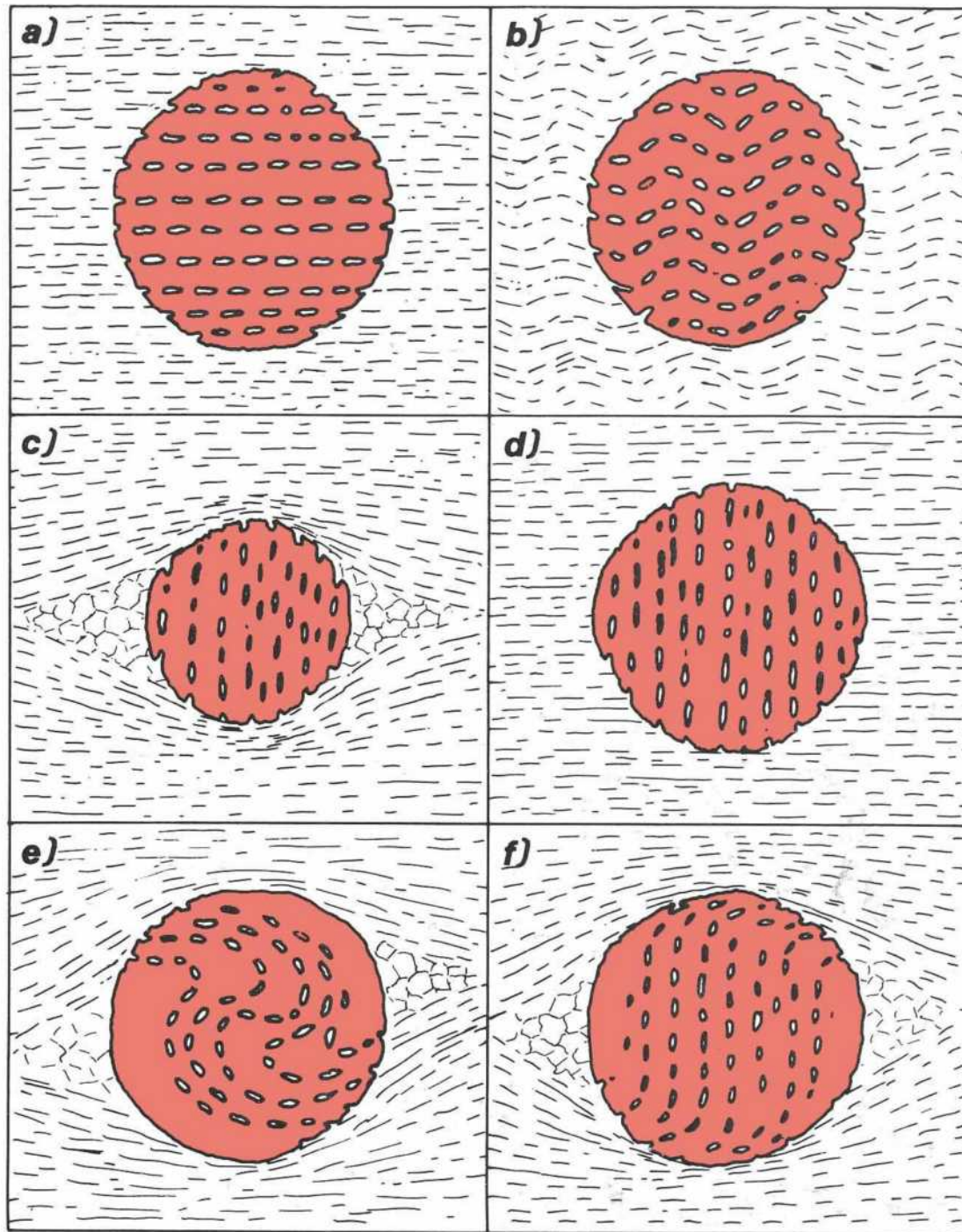


# Post-kinematic crystals

- a.** Helicitic folds **b.** Randomly oriented crystals **c.** Polygonal arcs **d.** Chiasmolite **e.** Late, inclusion-free rim on a poikiloblast (?) **f.** Random aggregate pseudomorph







**Post-kinematic:**  $S_i$  is identical to and continuous with  $S_e$

**Pre-kinematic:** Porphyroblasts are post- $S_2$ .  $S_i$  is inherited from an earlier deformation.  $S_e$  is compressed about the porphyroblast in (c) and a pressure shadow develops.

**Syn-kinematic:** Rotational porphyroblasts in which  $S_i$  is continuous with  $S_e$  suggesting that deformation did not outlast porphyroblast growth.

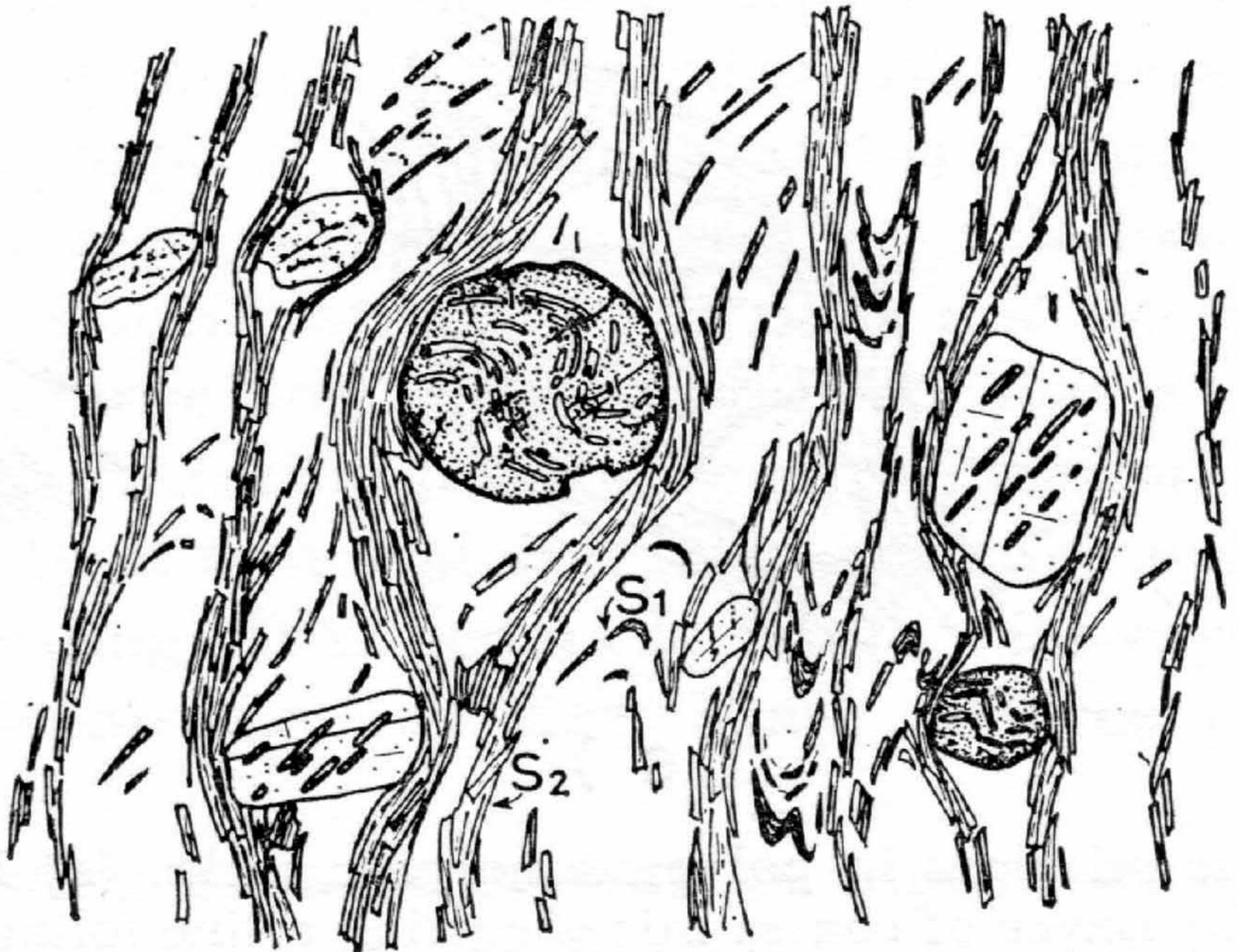
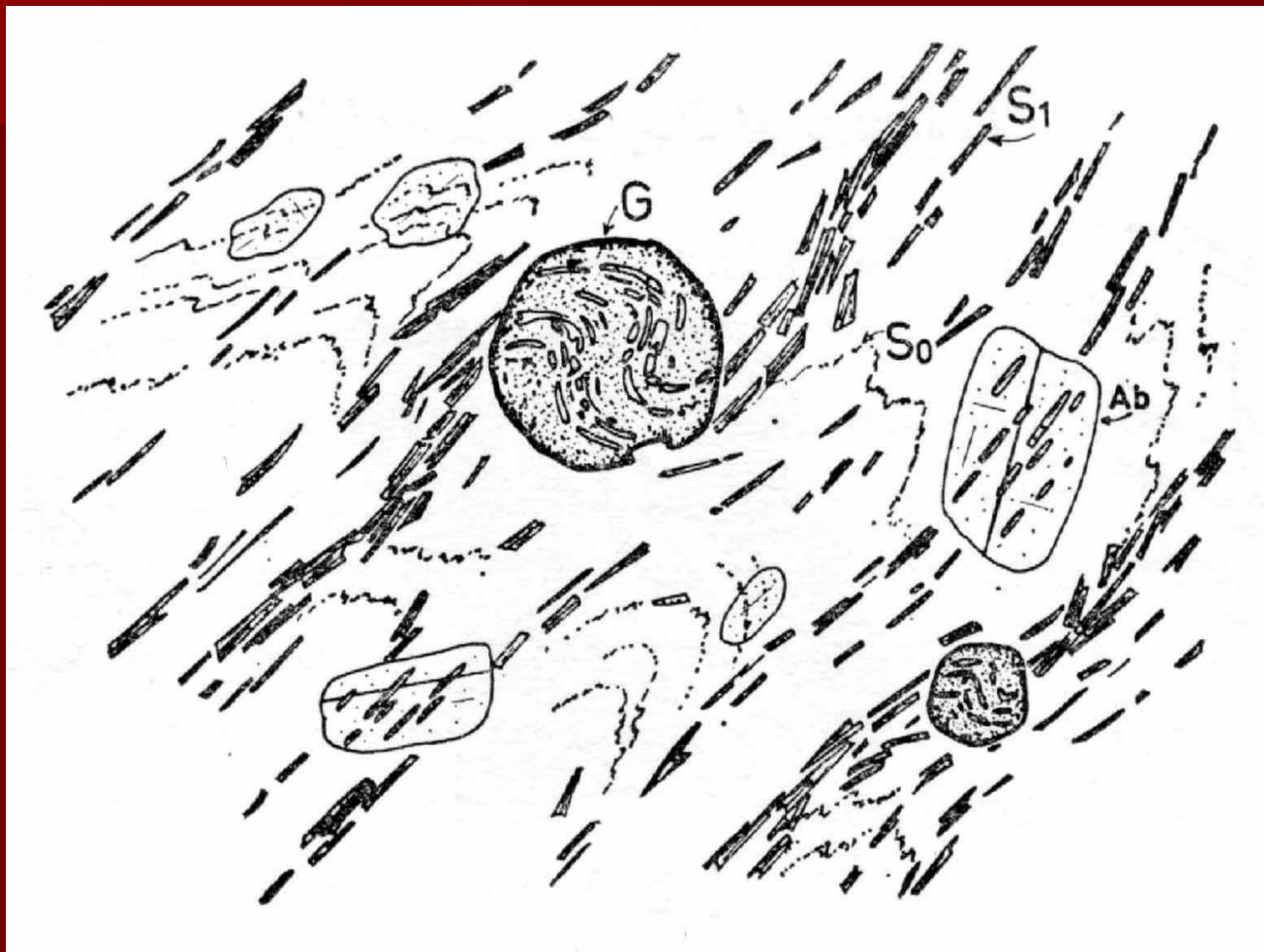


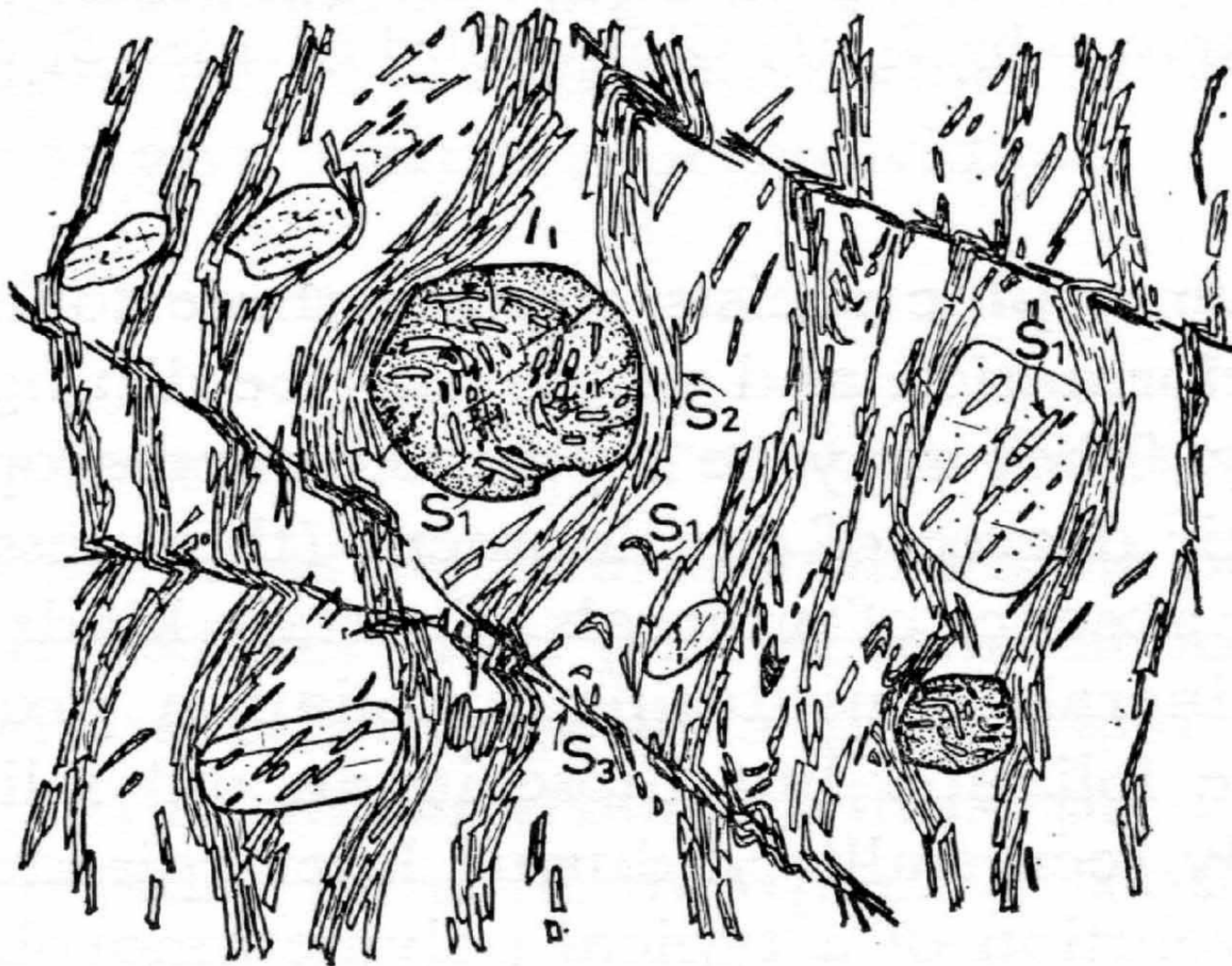
Figure 23.48b.



## Stages in Polymetamorphism ( $S_0$ , $S_1$ )



## Stages in Polymetamorphism ( $S_3$ )

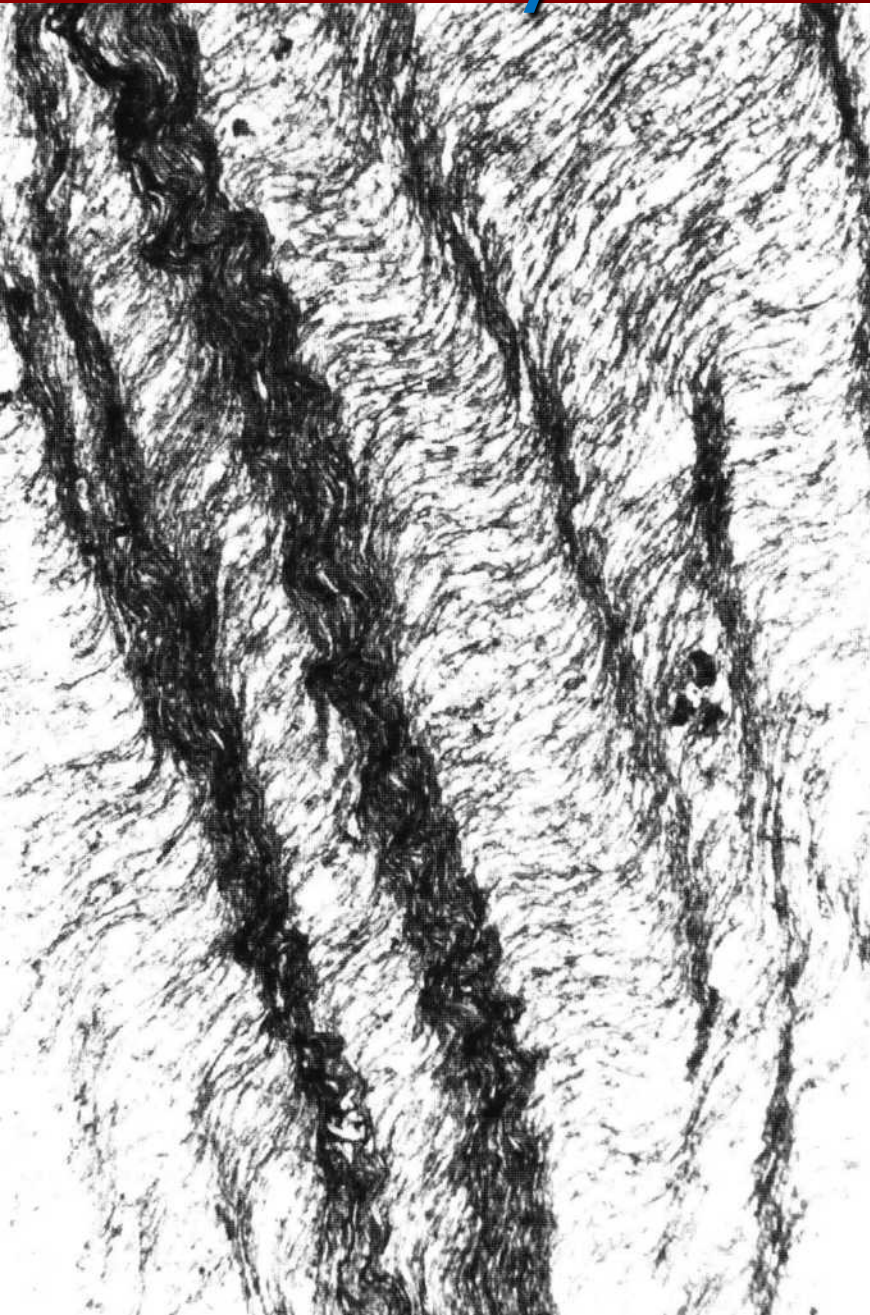




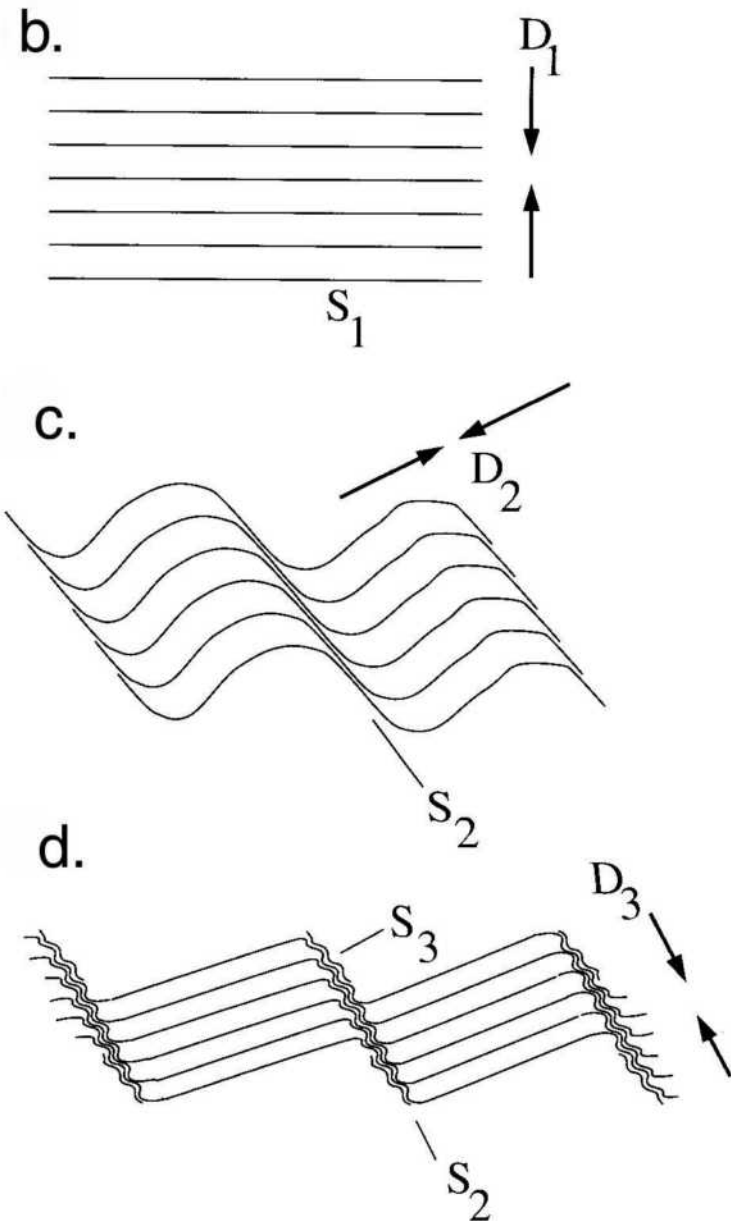
# Analysis of Deformed Rocks

- Deformational events:  $D_1$   $D_2$   $D_3$  ...
- Metamorphic events:  $M_1$   $M_2$   $M_3$  ...
- Foliations:  $S_0$   $S_1$   $S_2$   $S_3$  ...
- Lineations:  $L_0$   $L_1$   $L_2$   $L_3$  ...
- Plot on a metamorphism-deformation-time plot showing the crystallization of each mineral

# Analysis of Deformed Rocks

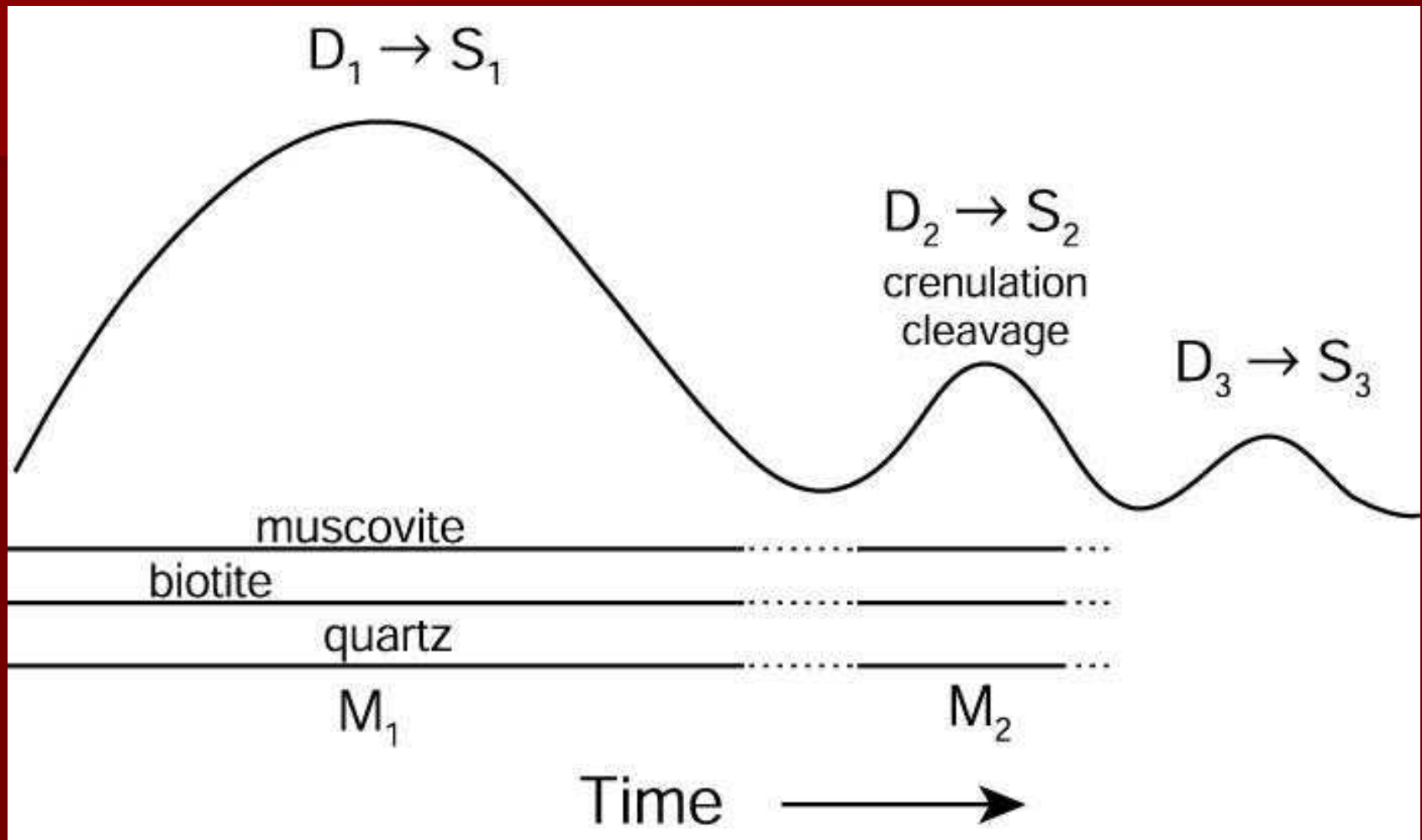


**Asymmetric crenulations cleavage ( $S_2$ ) developed over  $S_1$  cleavage.  $S_2$  is folded, as can be seen in the dark sub-vertical  $S_2$  bands.**





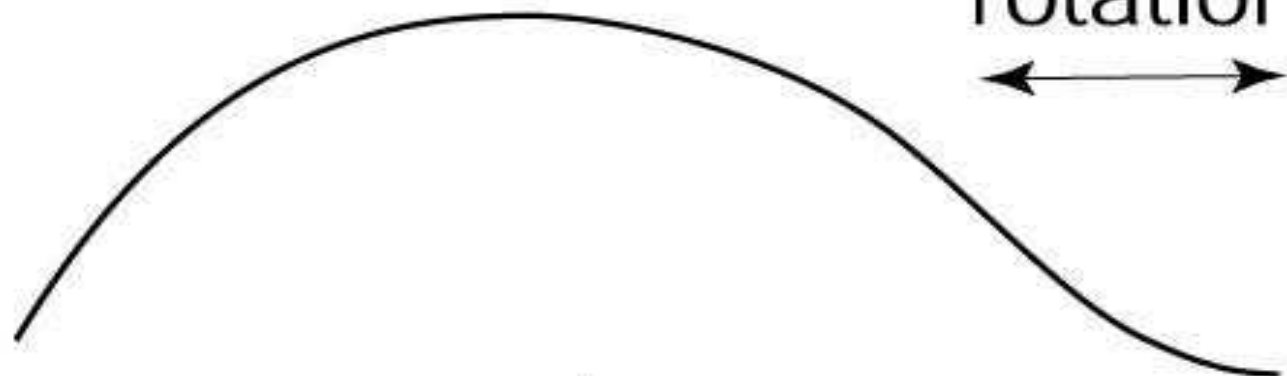
# Analysis of Deformed Rocks



Graphical analysis of the relationships between deformation (D), metamorphism (M),

$$D_1 \rightarrow S_1$$

rotation of  $S_1$



muscovite

biotite

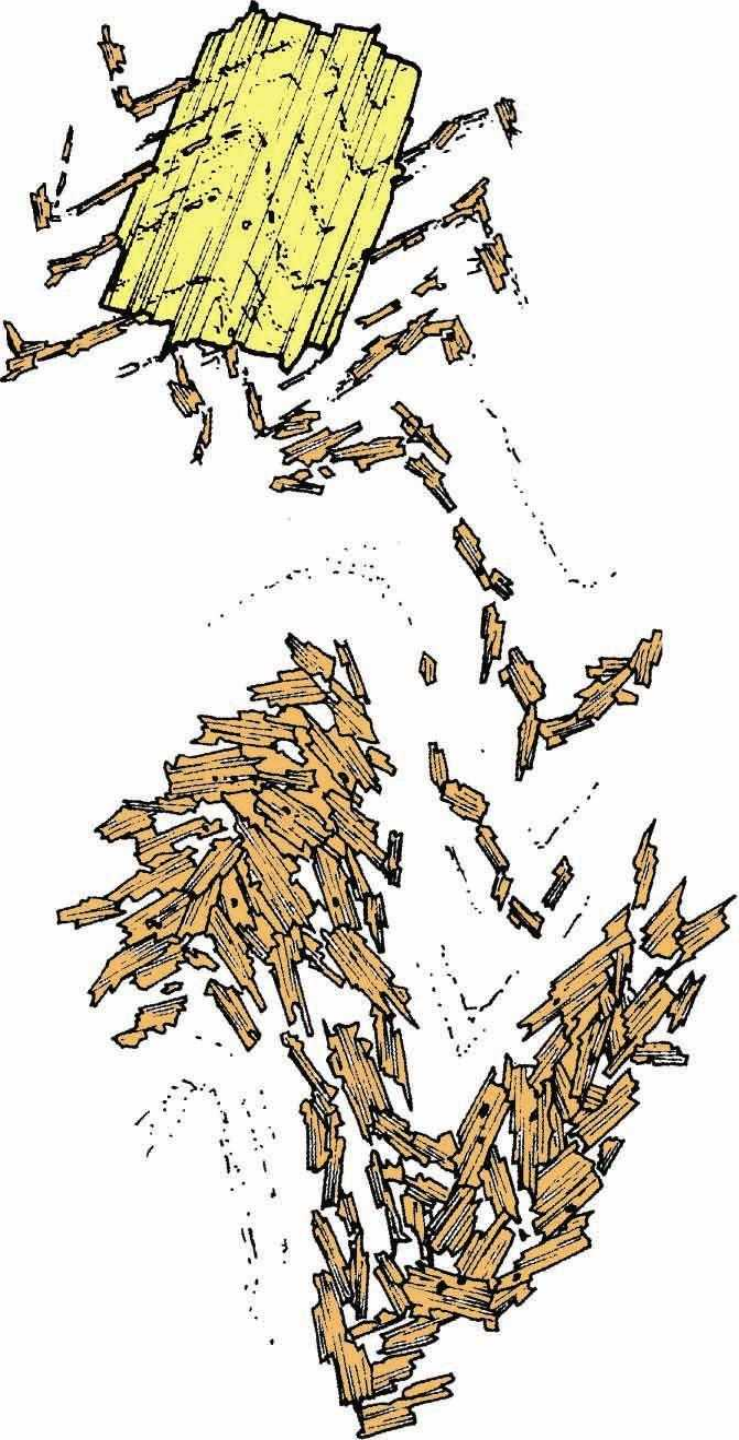
quartz

garnet

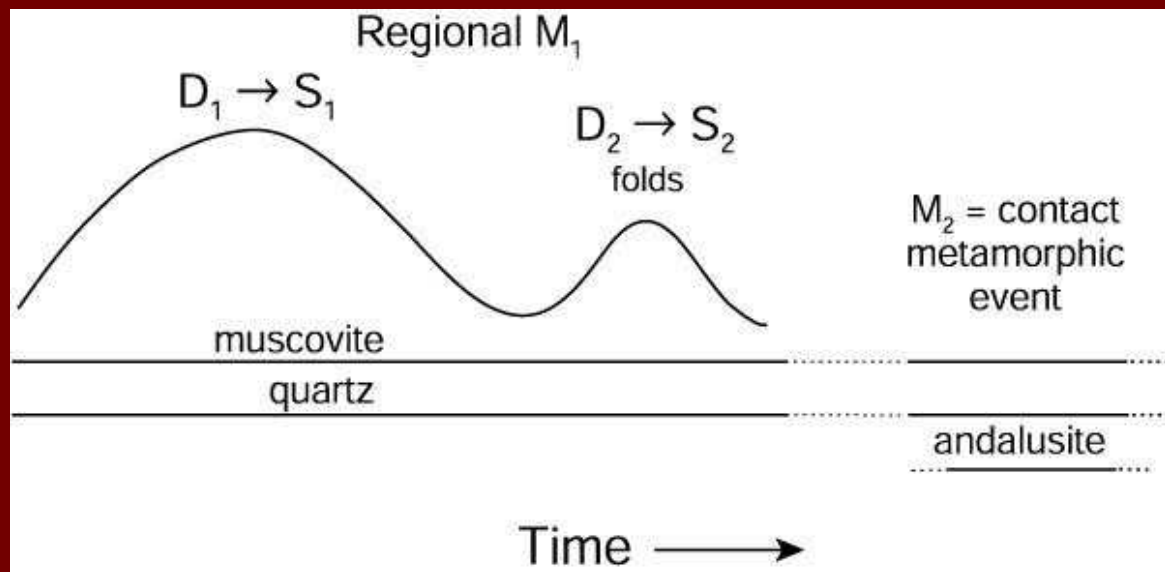
Time  $\longrightarrow$

Figure 23.45.





## Textures in a hypothetical andalusite porphyroblast-mica schist.



**Graphical analysis of the relationships between deformation (D), metamorphism (M), mineral growth**